

# Short course IV

## Thermochronology and landform development (Pecube-GUI)



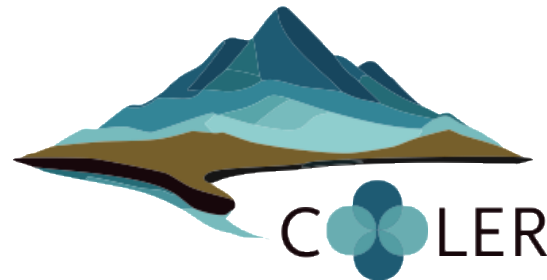
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with significant input from Maxime Bernard

*Now at Institute for Earth Surface Dynamics, Université de Lausanne, Switzerland*



Thermo2025 Kanazawa



# Short course presenters and contributors



Peter van der Beek



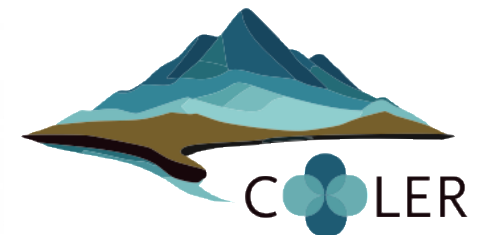
Lingxiao Gong



Isabel Wapenhans



Maxime Bernard

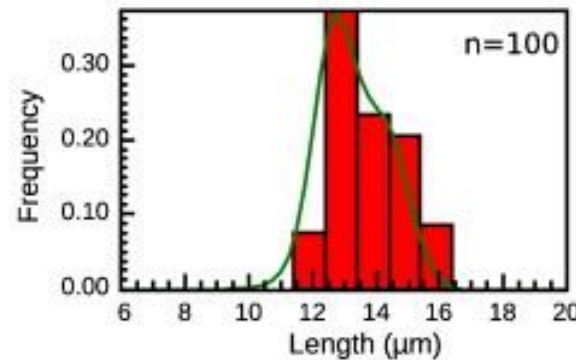


# Short course roadmap

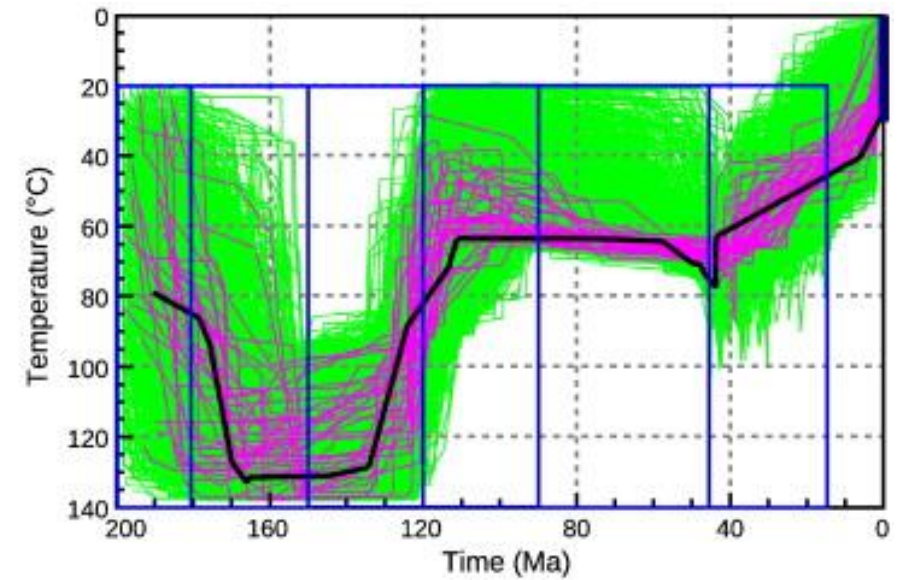
1. Pecube 1.01: Introduction to Pecube and Pecube-GUI (Peter)
2. Hands-on experience: running forward models in Pecube-GUI (Lingxiao)
3. Some notes on misfit (Peter)
4. Inverse modelling with Pecube / Pecube-GUI (Peter)
5. - if time permits – Inverse modelling exercise (Lingxiao)
6. Inverse modelling strategies and reporting inverse modelling results (Peter)

# Thermochronology $\Rightarrow$ thermal information

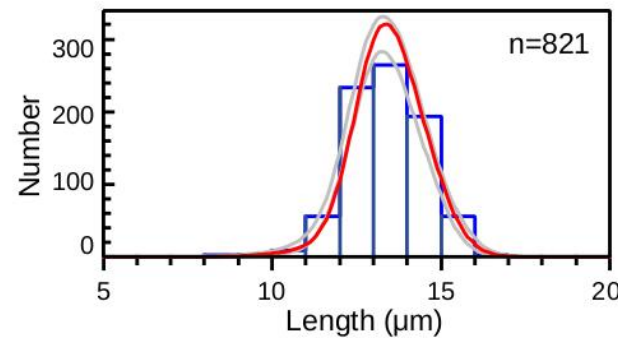
HeFTy



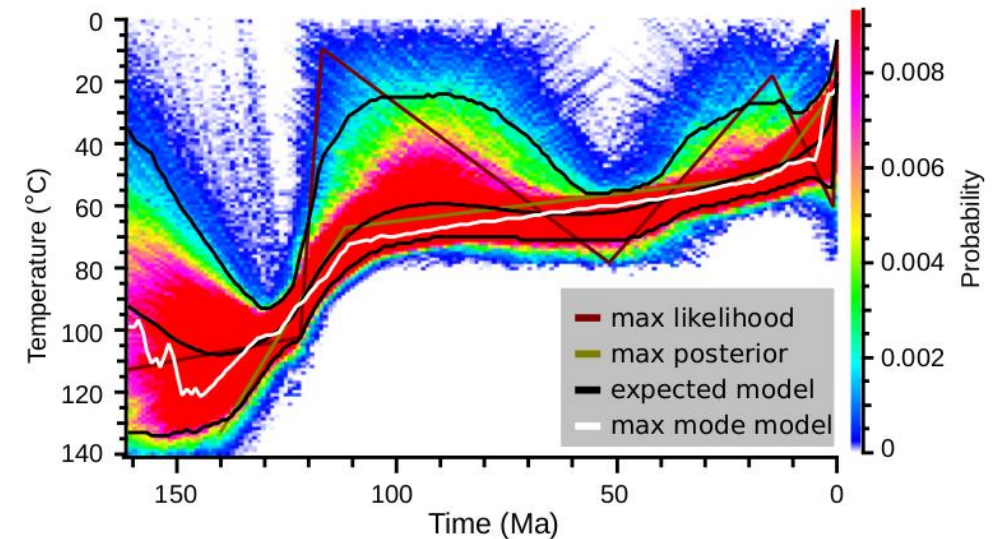
	MTL	AFT age	AHe age
Measured	$13.4 \pm 1.0$	$102 \pm 7$	$55 \pm 5$
Modelled	$13.3 \pm 1.1$	102	59



QTQt



	MTL	AFT age	AHe age
Measured	$13.4 \pm 1.2$	$102 \pm 7$	$55 \pm 5$
Modelled	$13.9 \pm 0.9$	$102 \pm 5$	$59 \pm 4$



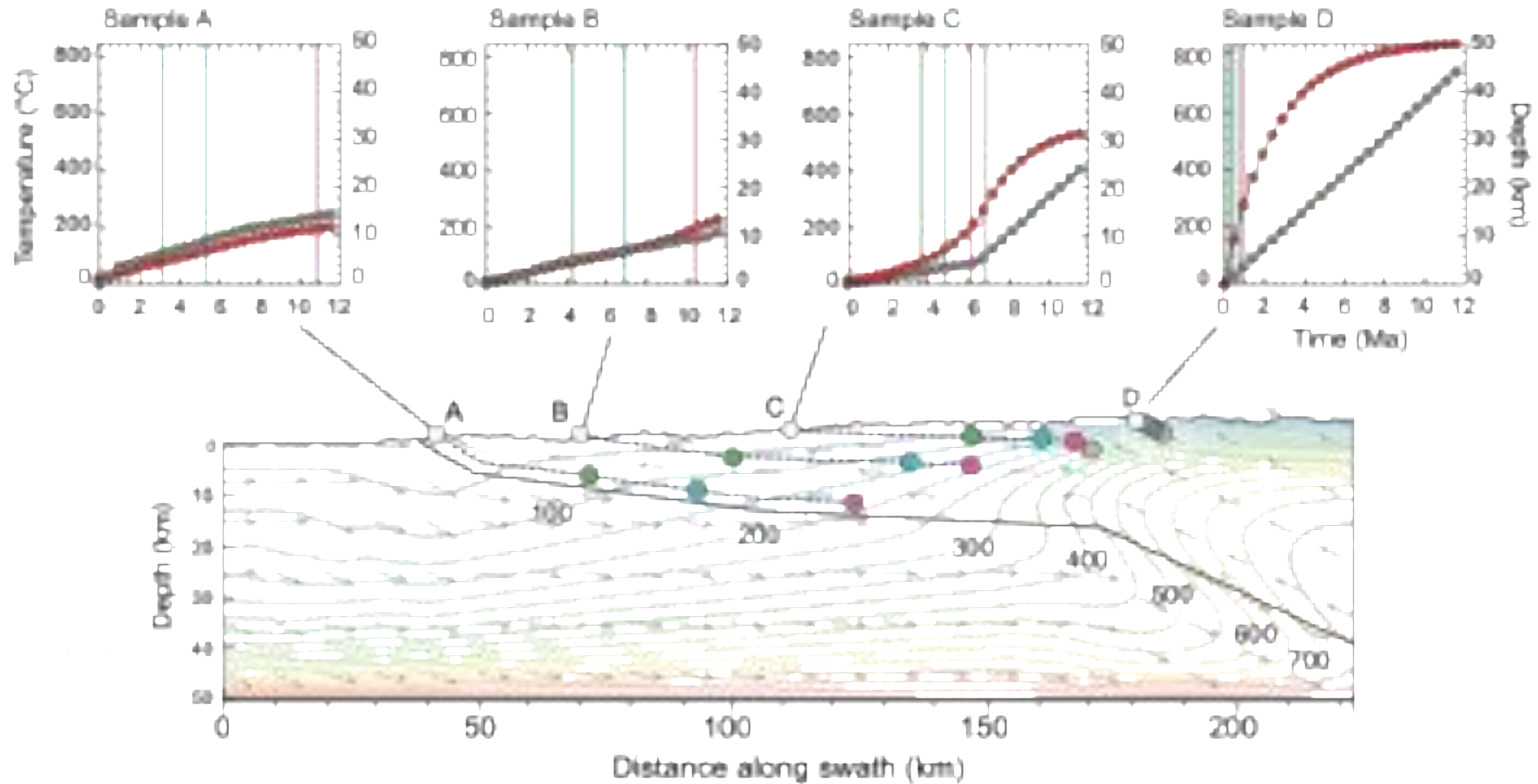
# Thermochronology $\Rightarrow$ thermal information

But we are interested in kinematic information!

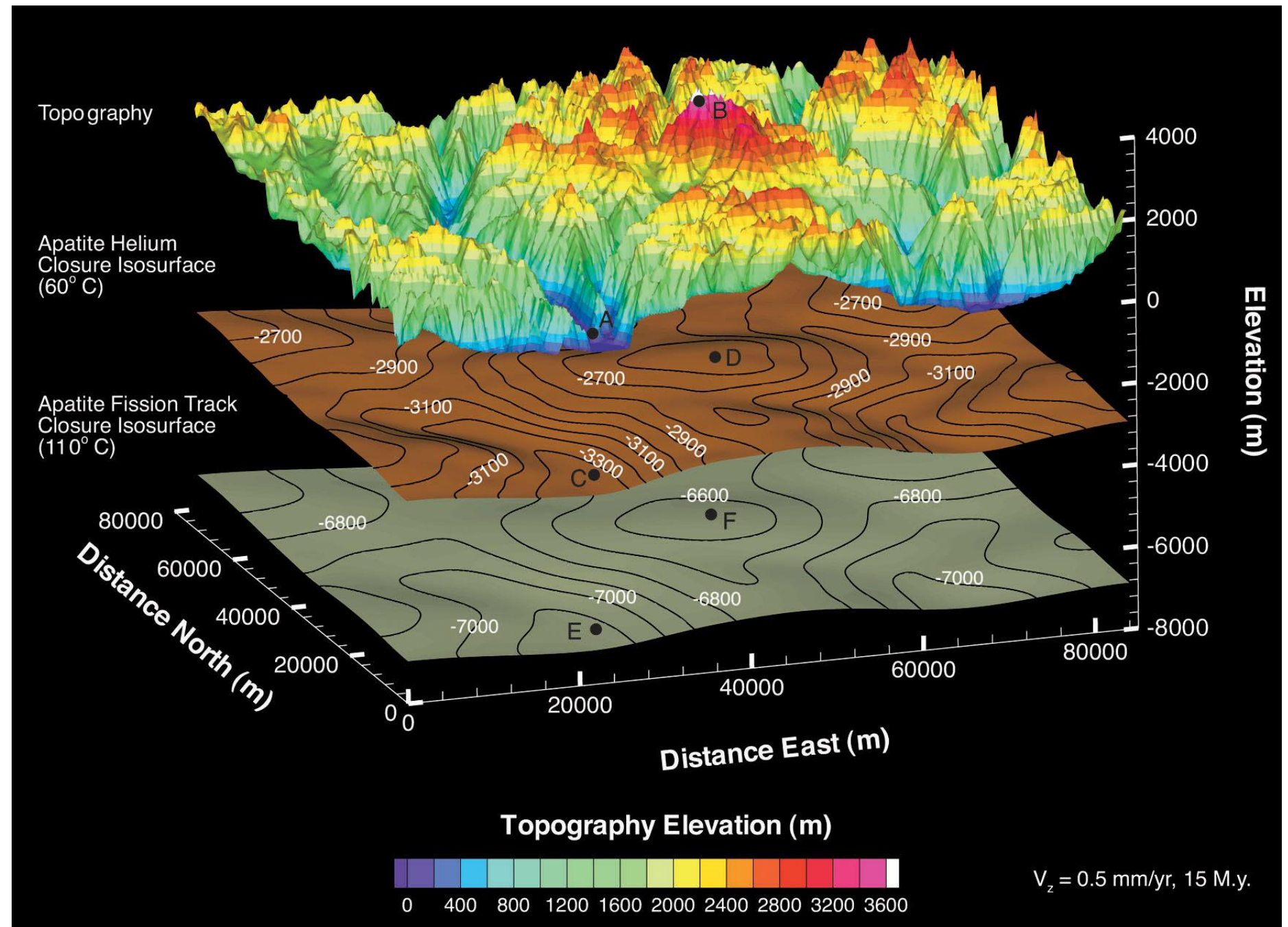
- When did exhumation occur, at what rate?
- At what time was this fault active?
- When did the current topography develop, what did it look like in the past?
- What does the pattern of thermochronology ages tell me about the tectonic history of this region?
- ...



# From thermal to kinematic reference frame

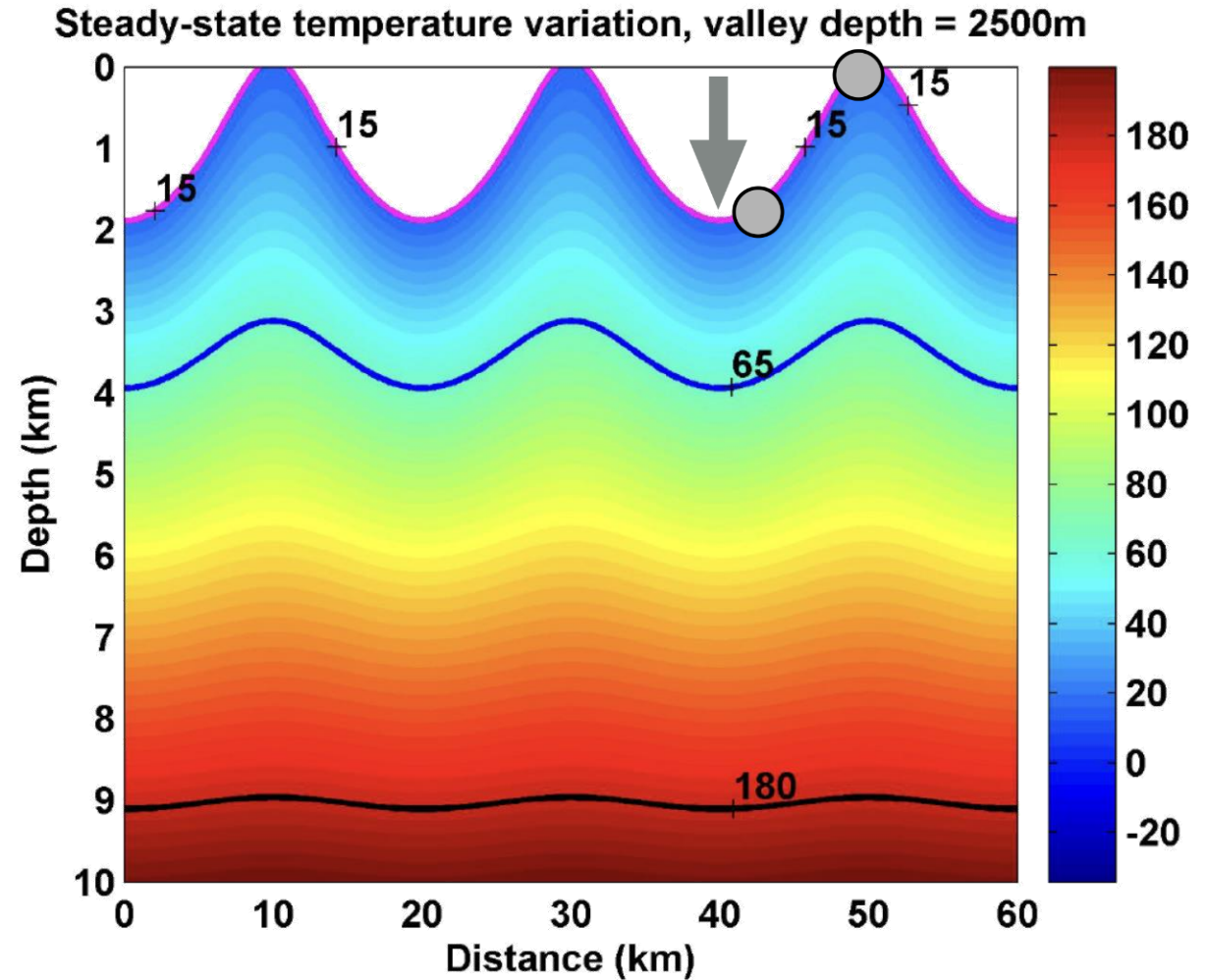
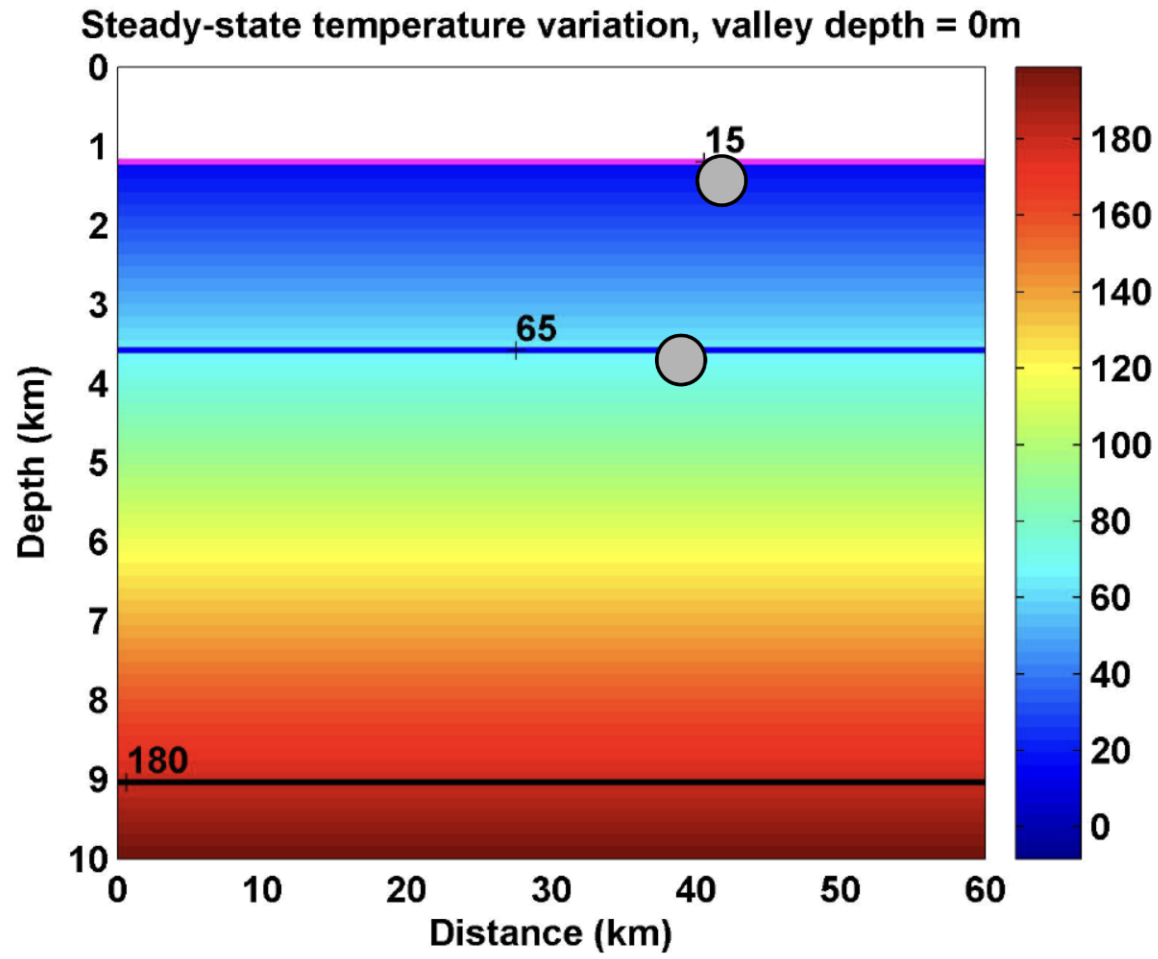


# Effect of topography on isotherms



Ehlers & Farley,  
EPSL 2003

# Using thermochronology to derive exhumation and relief histories





# Using thermochronology to derive exhumation and relief histories

⇒ solving the 3D transient heat-transport equation

$$\frac{dT}{dt} = \underbrace{\kappa \nabla^2 T}_{\text{Heat conduction}} + \underbrace{v \nabla T}_{\text{Heat advection}} + \underbrace{\frac{H}{\rho c}}_{\text{Heat production}}$$
$$\nabla T = \frac{\partial T}{\partial x} + \frac{\partial T}{\partial y} + \frac{\partial T}{\partial z}$$

Heat **conduction** (diffusion)    Heat **advection**    Heat **production** [K s<sup>-1</sup>]

$\kappa$  = Thermal diffusivity [m<sup>2</sup> s<sup>-1</sup> / km<sup>2</sup> My<sup>-1</sup>]

$\kappa = \frac{k}{\rho c}$  = Thermal conductivity [W m<sup>-1</sup> K<sup>-1</sup>]

$H$  = volumetric heat production [(μ)Wm<sup>-3</sup>]

$\rho$  = density [kg m<sup>-3</sup>]

$c$  = heat capacity [J kg<sup>-1</sup> K<sup>-1</sup>]

$v$  = rock/particle velocity [m s<sup>-1</sup>; km My<sup>-1</sup>]

$v_{(x,y,z,t)} = v_{x(t)} + v_{y(t)} + v_{z(t)}$

*Tectonics*

# Heat transport by conduction and advection

$$\frac{dT}{dt} = \kappa \nabla^2 T + v \nabla T + \frac{H}{\rho c}$$

**Characteristic timescales**

$L$ : “characteristic length” of system

conduction :  $\tau_c = \frac{L^2}{\kappa}$

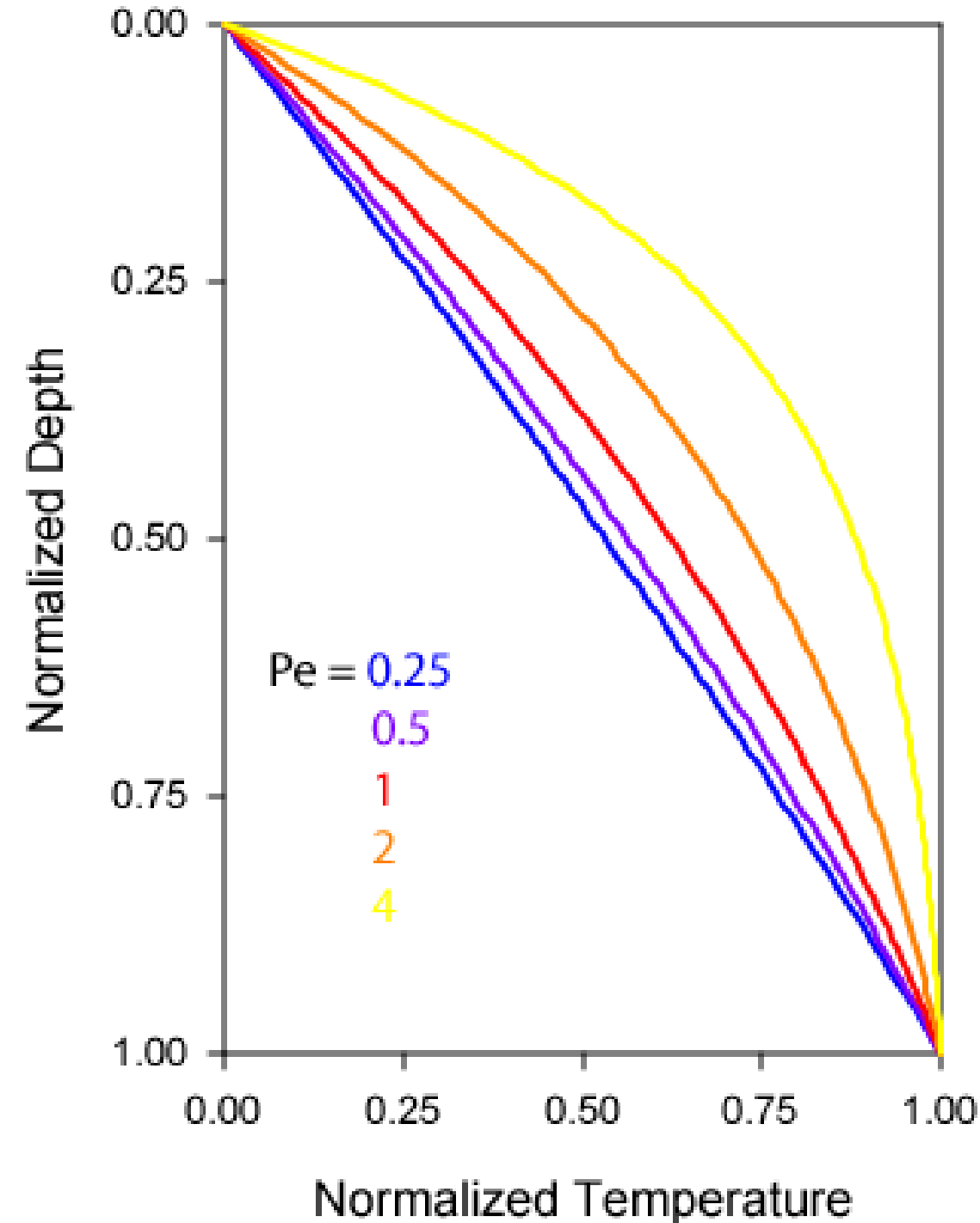
advection :  $\tau_a = \frac{L}{v}$

**Péclet number:**

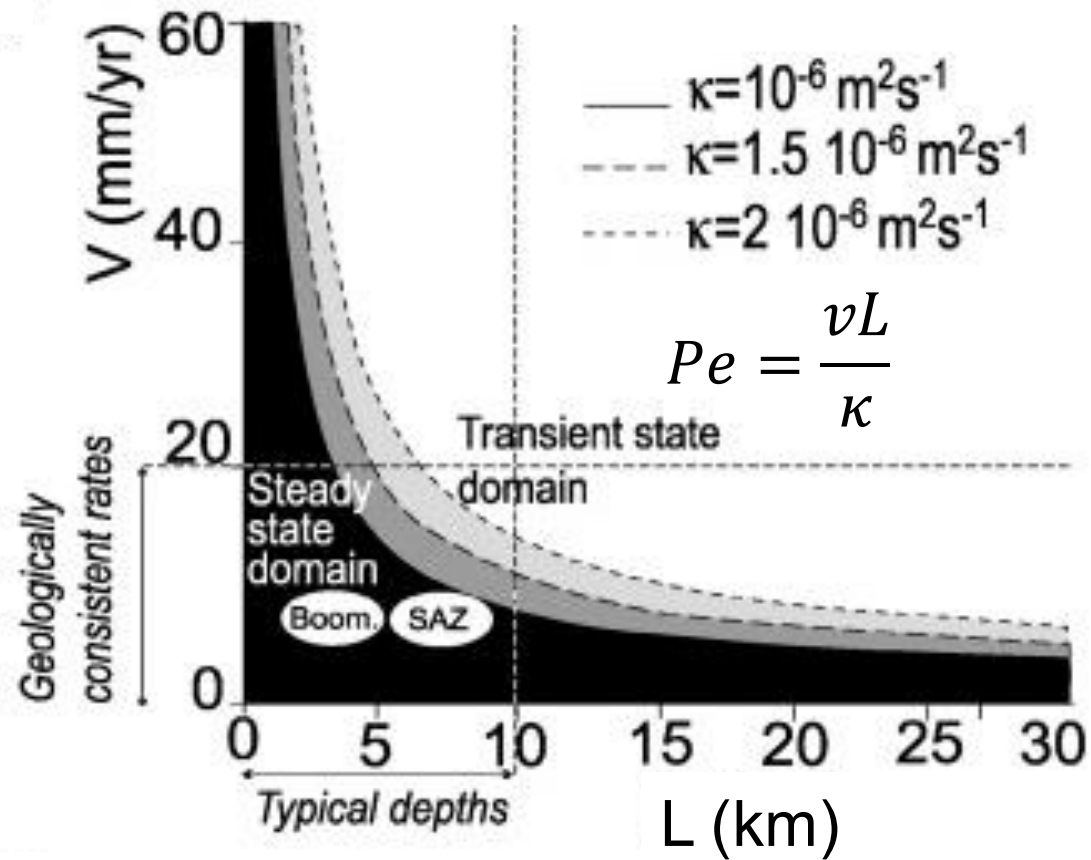
$$Pe = \frac{\tau_c}{\tau_a} = \frac{vL}{\kappa}$$

# Analytical steady-state solution to 1D heat conduction-advection equation

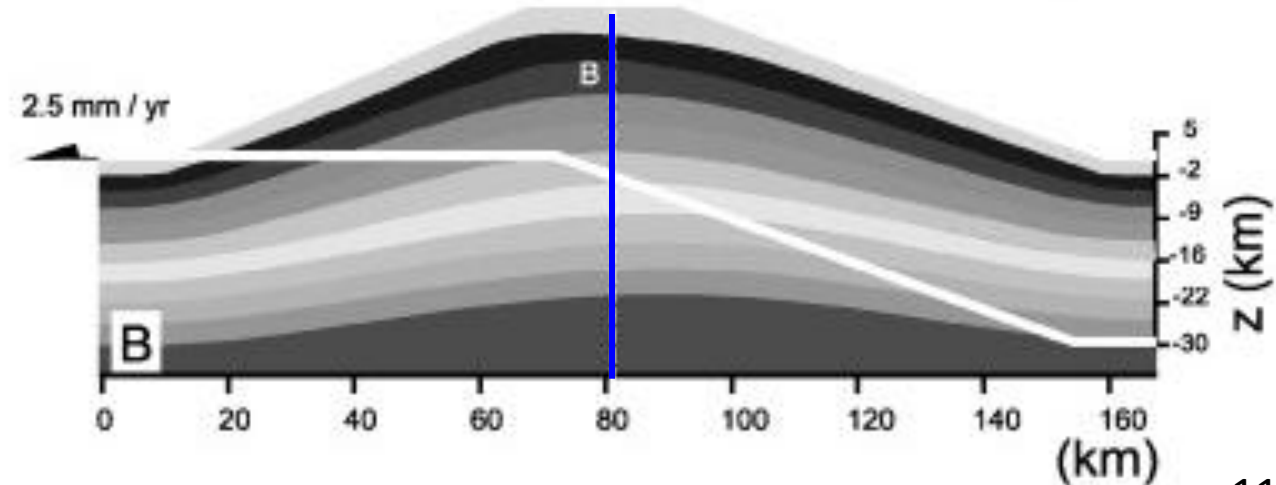
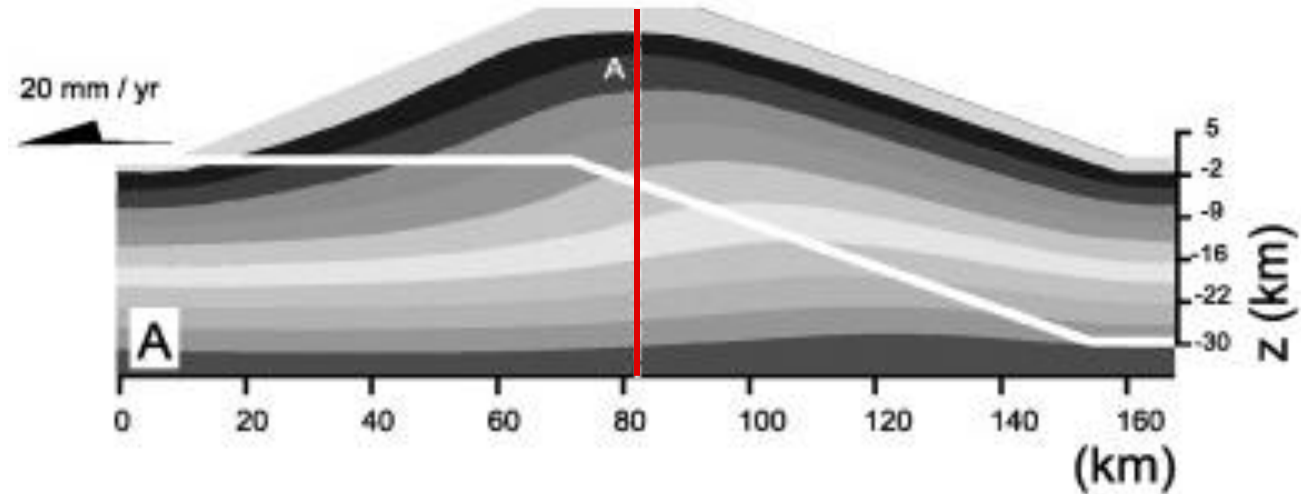
$$T'_{(z')} = \frac{(1 - e^{-Pe z'})}{(1 - e^{-Pe})}$$



# 2D heat advection-conduction example: Thrusting



$$Pe = \frac{vL}{\kappa}$$





# Using thermochronology to derive exhumation and relief histories

⇒ solving the 3D transient heat-transport equation

$$\frac{dT}{dt} = \kappa \nabla^2 T + v \nabla T + \frac{H}{\rho c}$$

$$\nabla T = \frac{\partial T}{\partial x} + \frac{\partial T}{\partial y} + \frac{\partial T}{\partial z}$$

Boundary conditions:

$$\rightarrow T_{(z=L)} = T_L$$

$$\rightarrow T_0 = \underbrace{T_{(z=0)}}_{\text{Sea-level temperature}} - \underbrace{G}_{\text{Surface temperature gradient}} \underbrace{h_{(x,y,t)}}_{\substack{\text{Surface evolution:} \\ \textit{Geomorphic history}}}$$

$$E_{(x,y,t)} = \underbrace{\int_0^t v_z(x,y,z,t)}_{\substack{\text{Exhumation} = \\ \text{Rock uplift} \\ \textit{Tectonics}}} - \underbrace{\Delta h_{(x,y,t)}}_{\substack{\text{Surface change} \\ \textit{Geomorphology}}}$$

# What Pecube does

## Input

- present-day topography (DEM)
- tectonic scenario ( $v_{(x, y, z, t)}$ )
- geomorphic scenario ( $h_{(x, y, t)}$ )
- (measured thermochronology data)

## Output

- Predicted thermochronology data (ages, ...) at surface; throughout model and/or at specific locations
- (misfit w.r.t. measured thermochronology data)

- 
- ```
graph TD; Input[Input] --> Process[Process]; Process --> Output[Output]
```
- Run through tectonic/geomorphic scenario to calculate initial position of points ending up at the surface
  - Solve transient heat equation by time-stepping finite-element solution
  - Track temperature/time paths for particles ending up at the surface
  - Use kinetic models to predict thermochronology data
  - (calculate misfit with measured thermochronology data)

## When is Pecube useful?

- When you have rapid (and variable) exhumation rates so that  $Pe > 1$  – for at least part of the history
- When you have lateral material transport
- When you have significant (and relatively rapid) changes in topography

# Pecube v1



PERGAMON

Computers & Geosciences 29 (2003) 787–794

COMPUTERS  
GEOSCIENCES

www.elsevier.com/locate/cageo

## Pecube: a new finite-element code to solve the 3D heat transport equation including the effects of a time-varying, finite amplitude surface topography<sup>☆</sup>

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Received 24 September 2001; received in revised form 30 September 2002; accepted 22 October 2002

### Abstract

A robust finite-element code (**Pecube**) has been developed to solve the three-dimensional heat transport equation in a crustal/lithospheric block undergoing uplift and surface erosion, and characterized by an evolving, finite-amplitude surface topography. The time derivative of the temperature field is approximated by a second-order accurate, mid-point, implicit scheme that takes into account the changing geometry of the problem. The method is based on a mixed Eulerian–Lagrangian approach that requires frequent re-interpolation of the temperature field in the vertical direction to ensure accuracy. From the computed crustal thermal structure, the temperature history of rock particles that, following an imposed tectonic scenario, are exhumed at the Earth's surface, is derived. These  $T - t$  paths can then be used to compute apparent isotopic ages for a range of geochronometers. The usefulness of the code is demonstrated by computing the predicted distribution of (U–Th)/He apatite ages in a high relief area of the Sierra Nevada, California, for a range of tectonic scenarios and comparing them to existing data.

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**Keywords:** Heat transfer equation; Numerical modelling; Time-varying surface topography; Low- $T$  geochronology; Relief evolution

### 1. Introduction

Much effort has recently been devoted to understanding the coupling between tectonics and erosion (Beaumont et al., 1999). Determining the rate at which landforms adapt to a changing tectonic environment has become an important problem to address. To achieve this, new low-temperature thermochronometric techniques have been developed such as (U–Th)/He apatite thermochronometry (Zaier et al., 1987; Farley et al., 1996; Wolf et al., 1996). To interpret age data from systems characterized by closure temperatures as low as 65°C, one must thus understand how surface

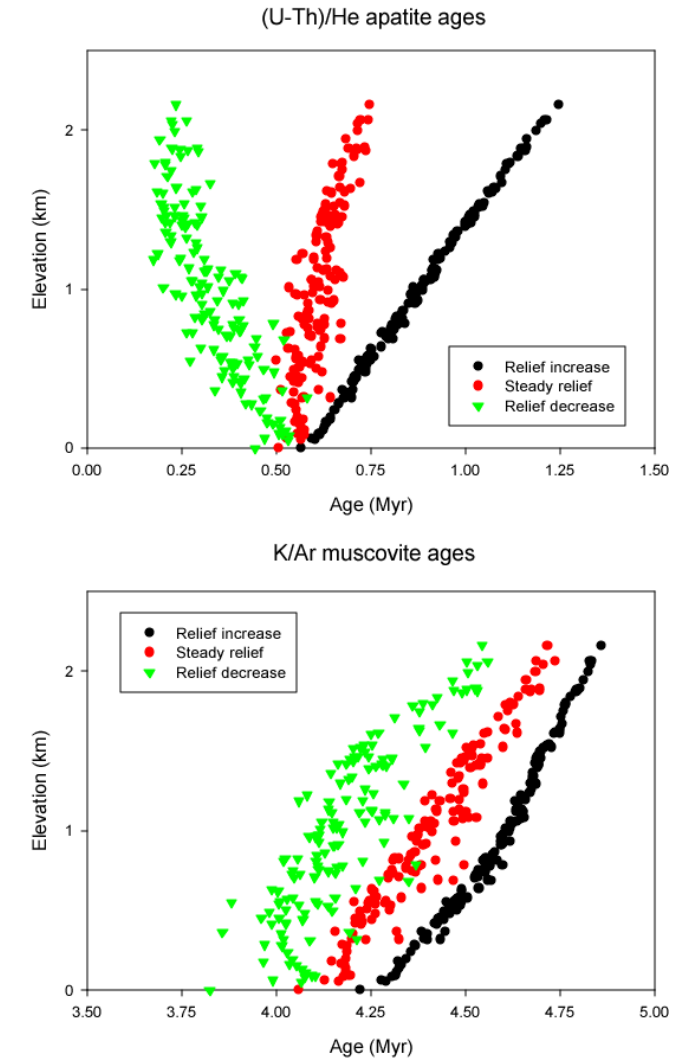
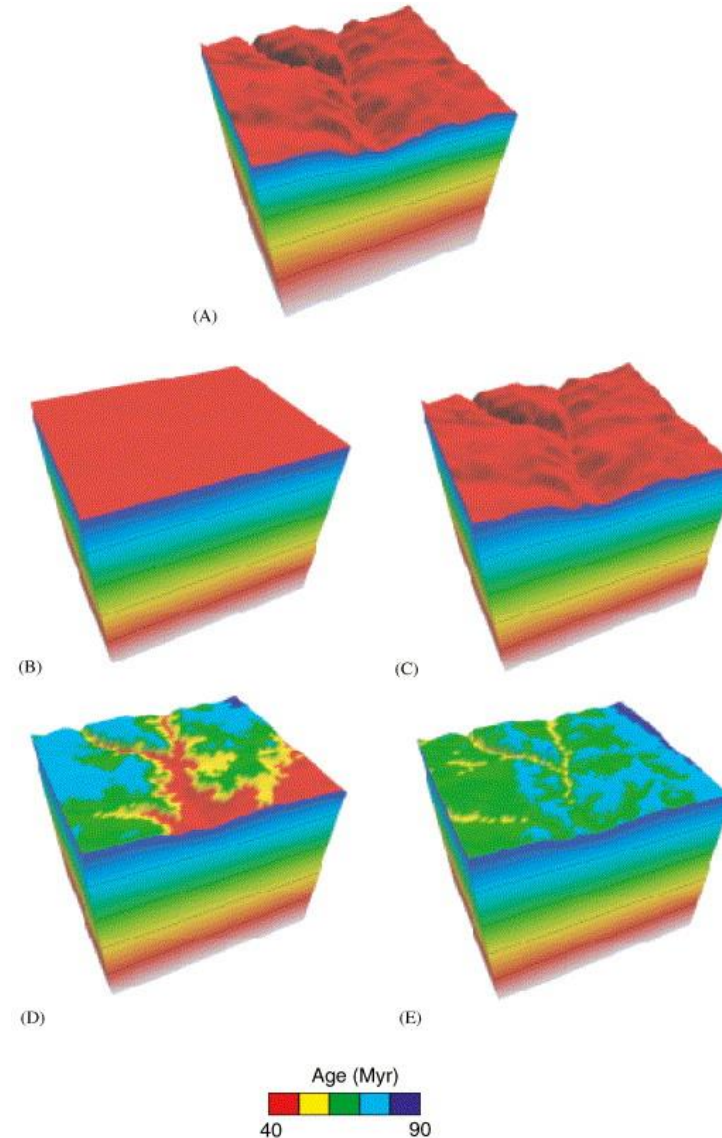
topography affects the thermal structure of the uppermost crust. Much work has already been devoted to this problem (Turcotte and Schubert, 1982; Stüwe et al., 1994; Mancktelow and Grasemann, 1997), but only the case of a “static” surface topography has been considered so far. To derive information on the rate of change of surface topography from thermochronometric data requires us, however, to better understand and quantify the effect of an evolving surface topography on the shape of the underlying isotherms.

Here, a new finite-element code is presented to solve the heat transfer equation in a crustal block undergoing uplift and denudation, and characterized by an evolving surface topography. The code has been designed for ease-of-use by non-specialists. It is available by contacting the author, or may be downloaded from the IAMG server.

<sup>☆</sup> Code available from server at <http://www.iang.org/CGEditor/index.htm>

<sup>\*</sup> Tel.: +61-2-6125-5512; fax: +61-2-6125-5443.

E-mail address: [jean.braun@anu.edu.au](mailto:jean.braun@anu.edu.au) (J. Braun).

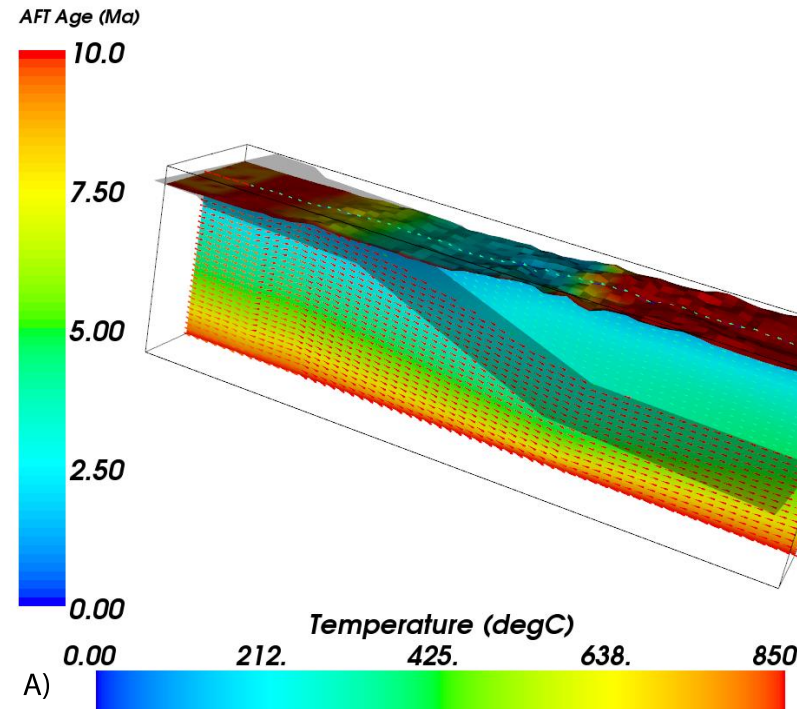




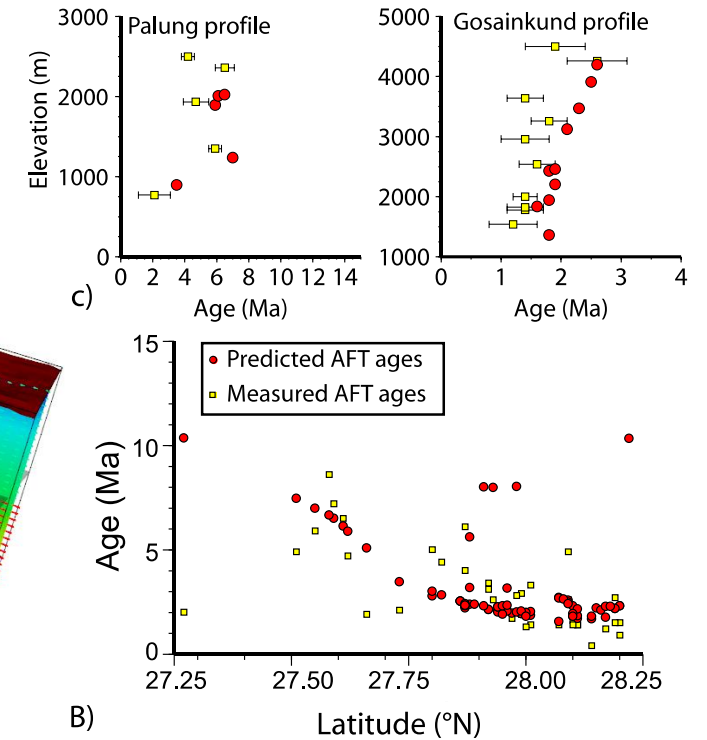
# Pecube v2



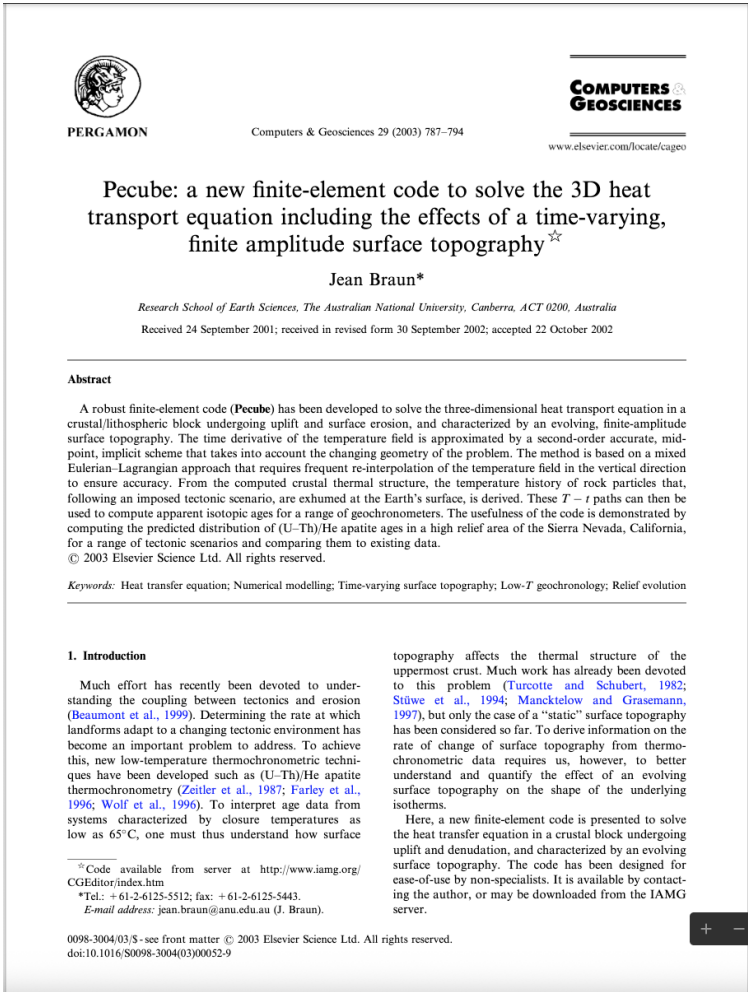
Braun et al., Tectonophysics 2012



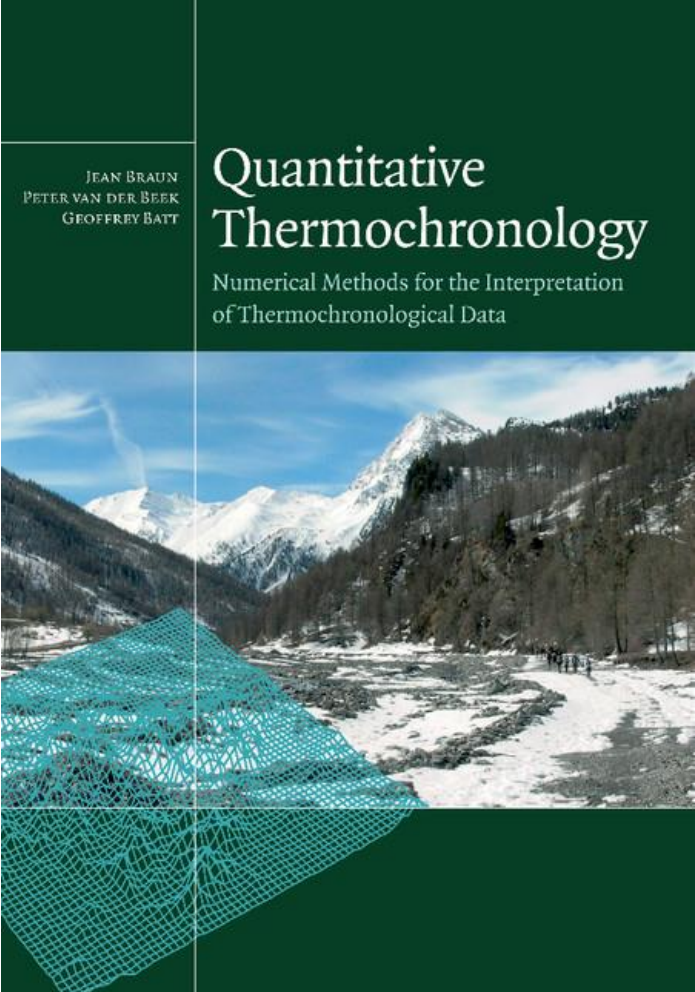
Robert et al., Geology 2009



# The Pecube “founding texts”



Braun, *Comp. Geosci.* 2003



Braun, van der Beek & Batt, *Quantitative Thermochronology*, CUP 2006



Braun et al., *Tectonophysics* 2012

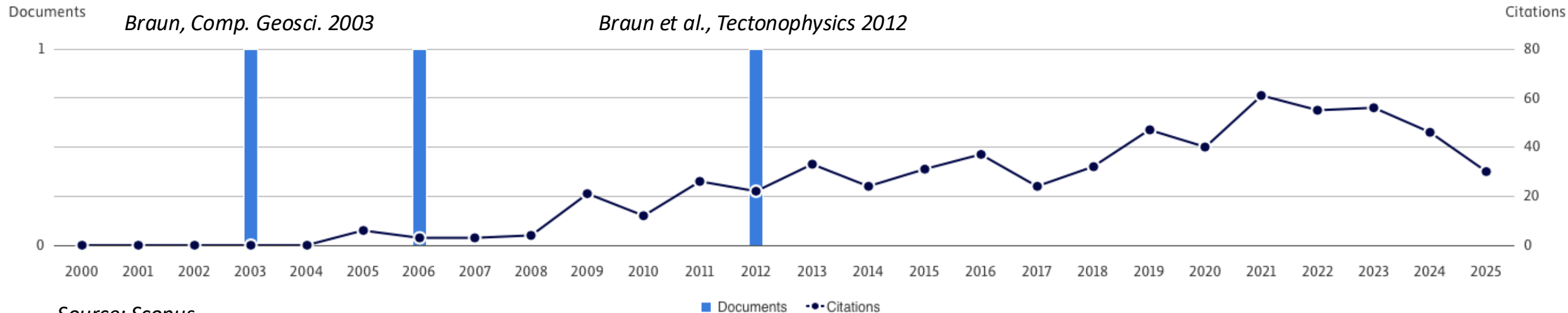
# The Pecube “founding father”





# Pecube citations

*Braun, van der Beek & Batt, Quantitative  
Thermochronology, CUP 2006*



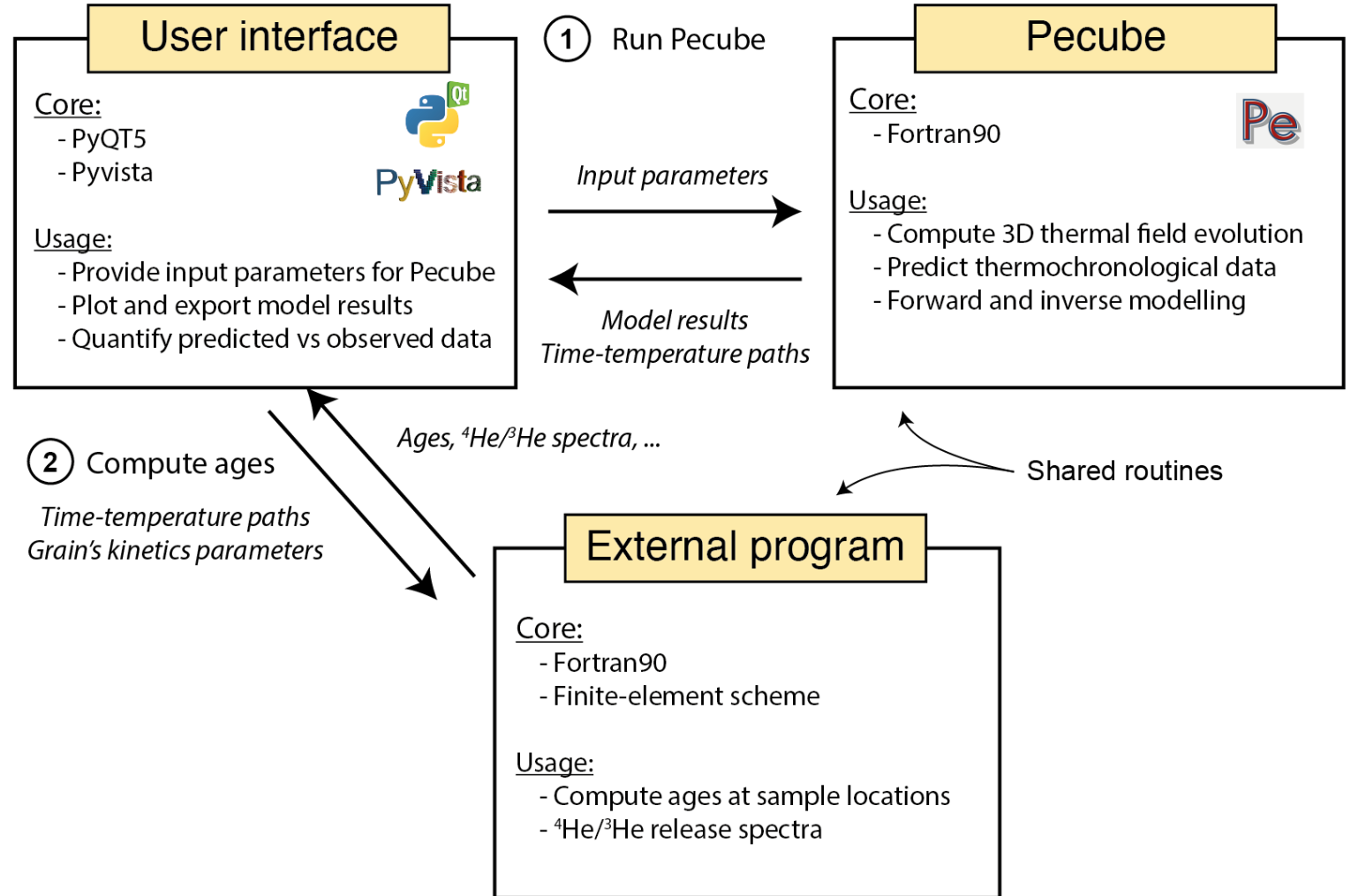
## Challenges:

- Not very user friendly (complex input-file structure, non-intuitive units for input parameters, no prior verification of model setup);
- Need to develop pre- and post-processing tools;
- No multi-kinetic age-prediction models; no possibility of obtaining sample-specific thermochronometric predictions.



# Pecube-GUI

- Python-based interface (PyQT5)
- Handles pre- and post-processing for Pecube
- Separated thermochronometric age predictions from Pecube core to provide more flexibility
- For Windows ( $\geq$  Windows 8) and MacOS ( $\geq$  11)
- Online documentation available



# Pecube-GUI versus most recent Pecube version

|                                             | Pecube v4 | PecubeGUI |
|---------------------------------------------|-----------|-----------|
| User interface                              | X         | V         |
| Sample-specific predictions                 | V         | V         |
| Grain-specific kinetics                     | X         | V         |
| Various kinetic models (AHe, ZHe, AFT, OSL) | V         | V         |
| Graphical output                            | X         | V         |
| Usage for teaching                          | X         | V         |

# Pecube-GUI website

## PecubeGUI

### What is PecubeGUI

PecubeGUI is a python-based program that uses PyQt5 to build the GUI. It is an open-source software distributed under the GNU license (soon released!)

The graphical user interface (GUI) aims to help any Pecube's user to set up their models with a user-friendly interface to provide the relevant input parameters. It also enables the users to directly plot the results within the interface, manipulate the characteristics of the plots, and export images. Finally, the PecubeGUI facilitates the coupling between Pecube and the glacial landscape evolution model iSOSIA (Egholm et al., 2011), by enabling the user to load iSOSIA output files, that can be read and used by Pecube. However, output files from other surface processes models can also be loaded.

PecubeGUI comes with a new version of Pecube that enables to predict (U-Th)/He ages on apatite for specific locations. It also includes some radiation damage models, alpha ejection distances, the possibility to set the characteristics of each grain (i.e., grain size, uranium and thorium concentration), and to predict 4He/3He profiles.

### Visualization tools for PecubeGUI

PyPIVoT (PecubeGUI Visualization Tools) are a collection of newly developed Python- and Matlab-based visualization scripts designed to help users of PecubeGUI better their understanding of inversion parameter interdependence and the parameter space explored in inversion. PecubeGUI is a thermal-kinematic modelling software used in thermochronology. [\[learn more\]](#)

### Download PecubeGUI

PecubeGU-beta is now available!  
Download from [Zenodo](#)

PecubeGUI works for MacOS > 11.0 (BigSur) and Windows. Please download the package that matches your chip (MacOS M1/M2 or Intel, Windows\_exe)

### PecubeGUI Documentation

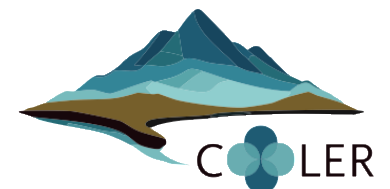
You can access the documentation, here: [PecubeGUI – Documentation](#)

### Example input files

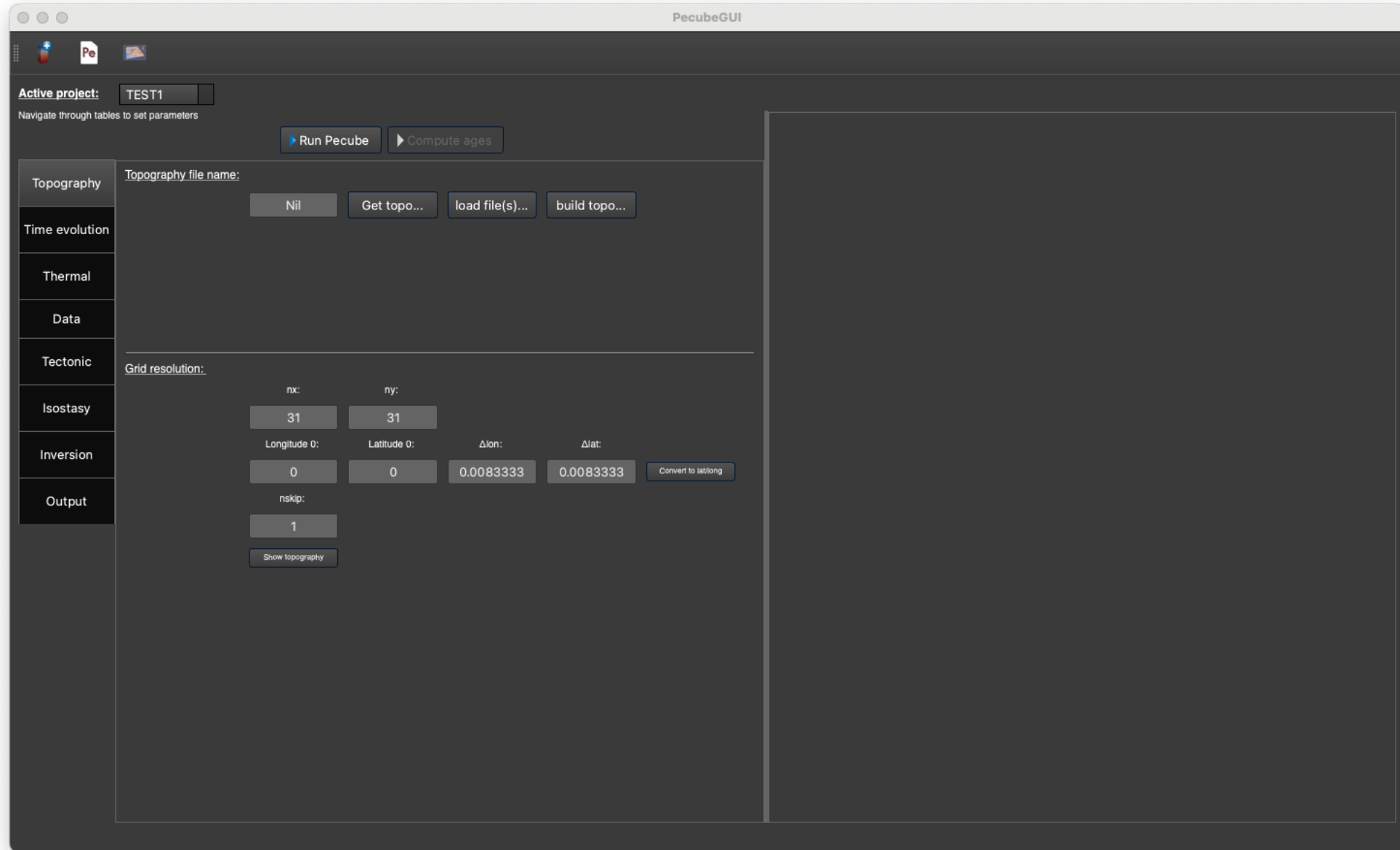
- [Forward modelling \(zip-file, 24 MB\)](#)
- [Inversion \(zip-file, 4,7 MB\)](#)



<https://erc-cooler.eu/pecubegui/>



# Let's try this!



# Misfit

There are MANY different misfit (*objective*) functions / definitions of misfit ...

The original version of Pecube used:  $\phi = \frac{1}{N} \sqrt{\sum_{i=1}^N \frac{(\alpha_{i,mod} - \alpha_{i,dat})^2}{\sigma_i^2}}$

Since Pecube v2 (2012) the  $\chi^2$ -statistic is used:

$$\phi = \sum_{i=1}^N \left( \frac{\alpha_{i,mod} - \alpha_{i,dat}}{\sigma_i} \right)^2$$

$\phi$  = misfit

$N$  = number of data

$\alpha_{i,mod}$  = modelled age of  $i^{\text{th}}$  datapoint

$\alpha_{i,dat}$  = observed age of  $i^{\text{th}}$  datapoint

$\sigma_i$  = uncertainty of  $i^{\text{th}}$  datapoint



# Misfit

Advantages of the  $\chi^2$ -statistic  $\phi = \sum_{i=1}^N \left( \frac{\alpha_{i,mod} - \alpha_{i,dat}}{\sigma_i} \right)^2$ :

$\phi$  = misfit

$N$  = number of data

$\alpha_{i,mod}$  = modelled age of  $i^{\text{th}}$  datapoint

$\alpha_{i,dat}$  = observed age of  $i^{\text{th}}$  datapoint

$\sigma_i$  = uncertainty of  $i^{\text{th}}$  datapoint

$p$  = number of free parameters

1. It can be easily transformed into the *reduced*  $\chi^2$ -statistic:

$$\phi_r = \frac{1}{N-p-1} \sum_{i=1}^N \frac{(\alpha_{i,mod} - \alpha_{i,dat})^2}{\sigma_i^2}$$

Which gives insight into how well the data are fit: if  $\phi_r < 1$  then all datapoints are fit to within  $1\sigma$  error (if  $\phi_r < 2$  then fit within  $2\sigma$ , etc.)

# Misfit

Advantages of the  $\chi^2$ -statistic  $\phi = \sum_{i=1}^N \left( \frac{\alpha_{i,mod} - \alpha_{i,dat}}{\sigma_i} \right)^2$ :

$\phi$  = misfit

$N$  = number of data

$\alpha_{i,mod}$  = modelled age of  $i^{\text{th}}$  datapoint

$\alpha_{i,dat}$  = observed age of  $i^{\text{th}}$  datapoint

$\sigma_i$  = uncertainty of  $i^{\text{th}}$  datapoint

$p$  = number of free parameters

1. It can be easily transformed into the *reduced*  $\chi^2$ -statistic.
2. We can use it to calculate the *log-likelihood*:

$$\log(L) = - \sum_{i=1}^N \left[ \frac{\ln(2\pi)}{2} + \ln(\sigma_i) + 0.5 \left( \frac{\alpha_{i,mod} - \alpha_{i,dat}}{\sigma_i} \right)^2 \right]$$

3. Which can be used to calculate the *Bayesian Information Criterion*:

$$BIC = -2 \log(L) + p \log(N)$$

The *BIC* informs us about the appropriate level of model complexity

# Misfit

Objective functions / what we try to achieve:

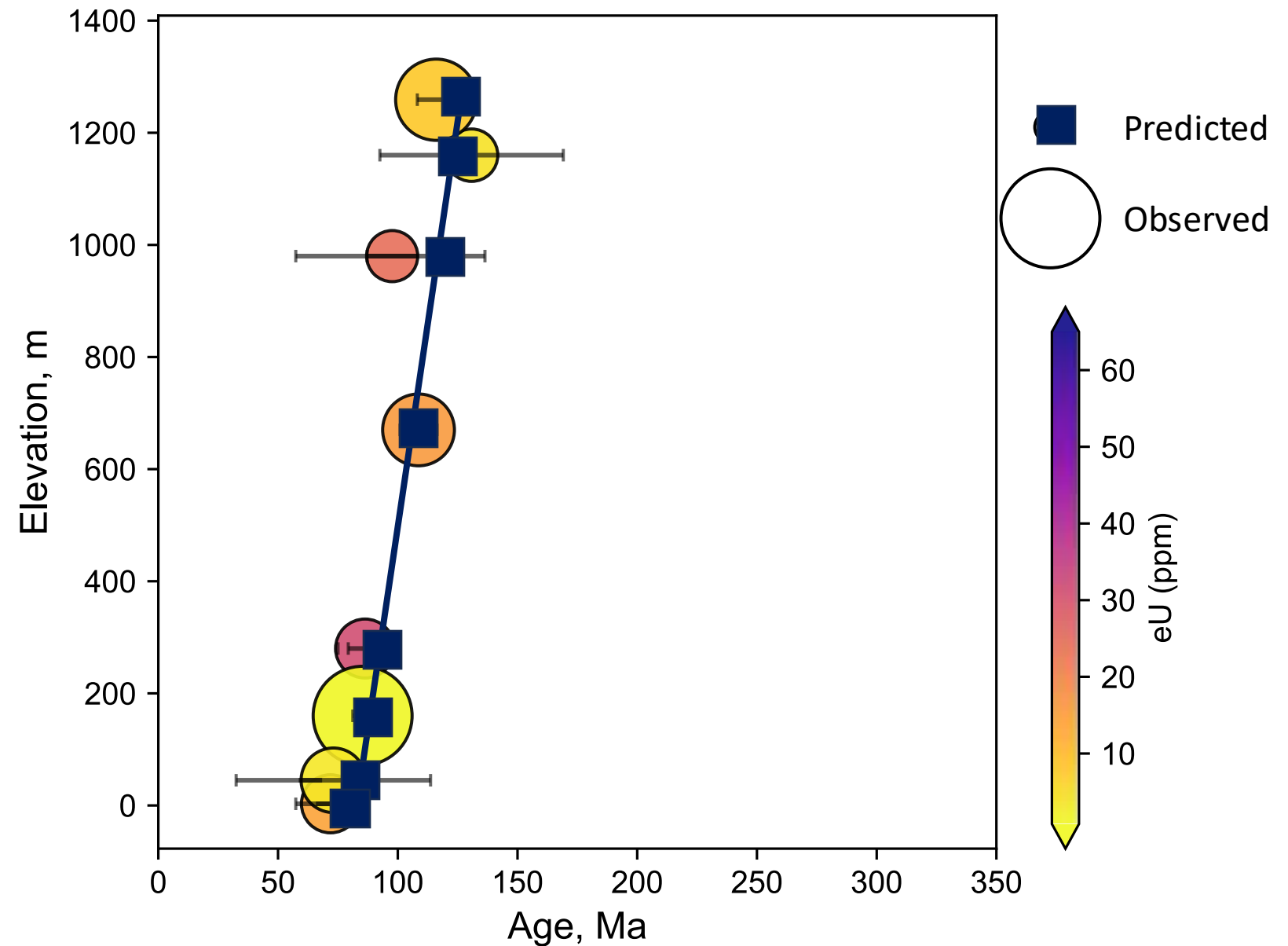
- *Minimise* the  $\chi^2$  / reduced  $\chi^2$
- *Maximise* the log-likelihood (note:  $\log(L) < 0$  by definition)
- *Minimise* the *BIC*

# Misfit

A simple case:

$\phi_r < 1$ ;  
 $\phi$  small;  
 $\log(L) \rightarrow 0$

Good fit to data



# Misfit

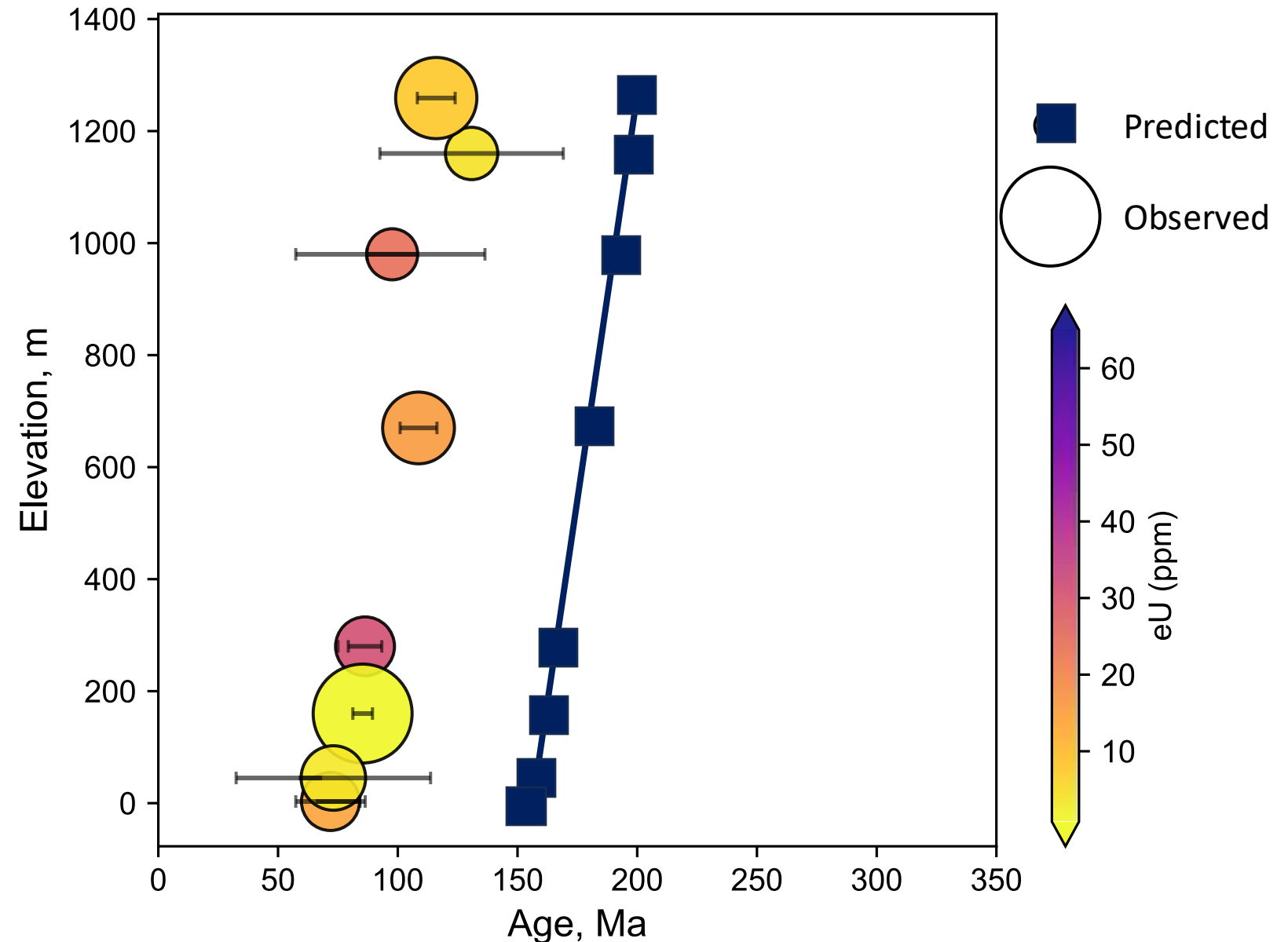
A simple case:

$$\phi_r > 1;$$

$\phi$  large;

$$\log(L) \ll 0$$

Bad fit to data

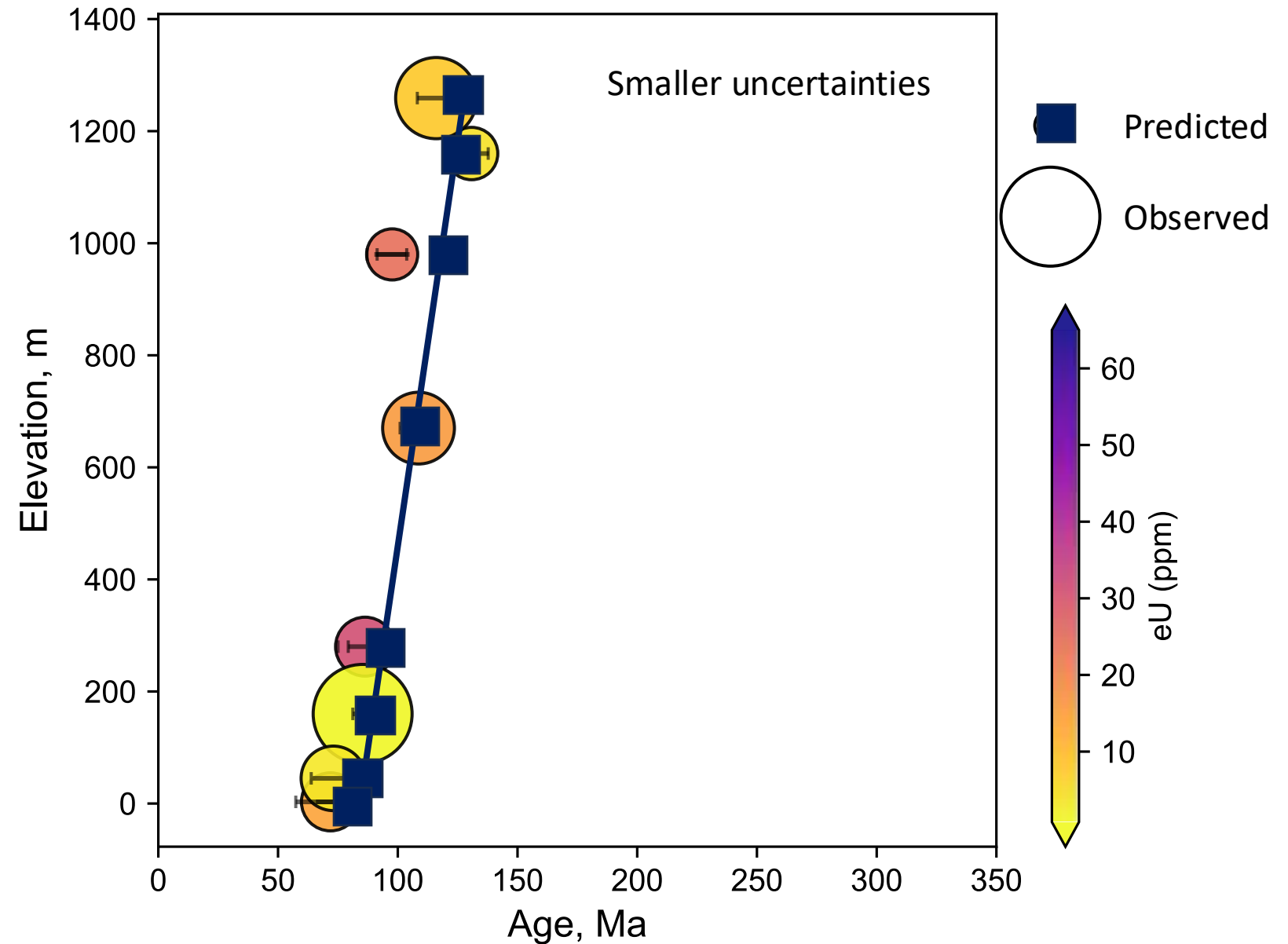


# Misfit

But:

$\phi_r > 1$ ;  
 $\phi$  larger;  
 $\log(L) < 0$

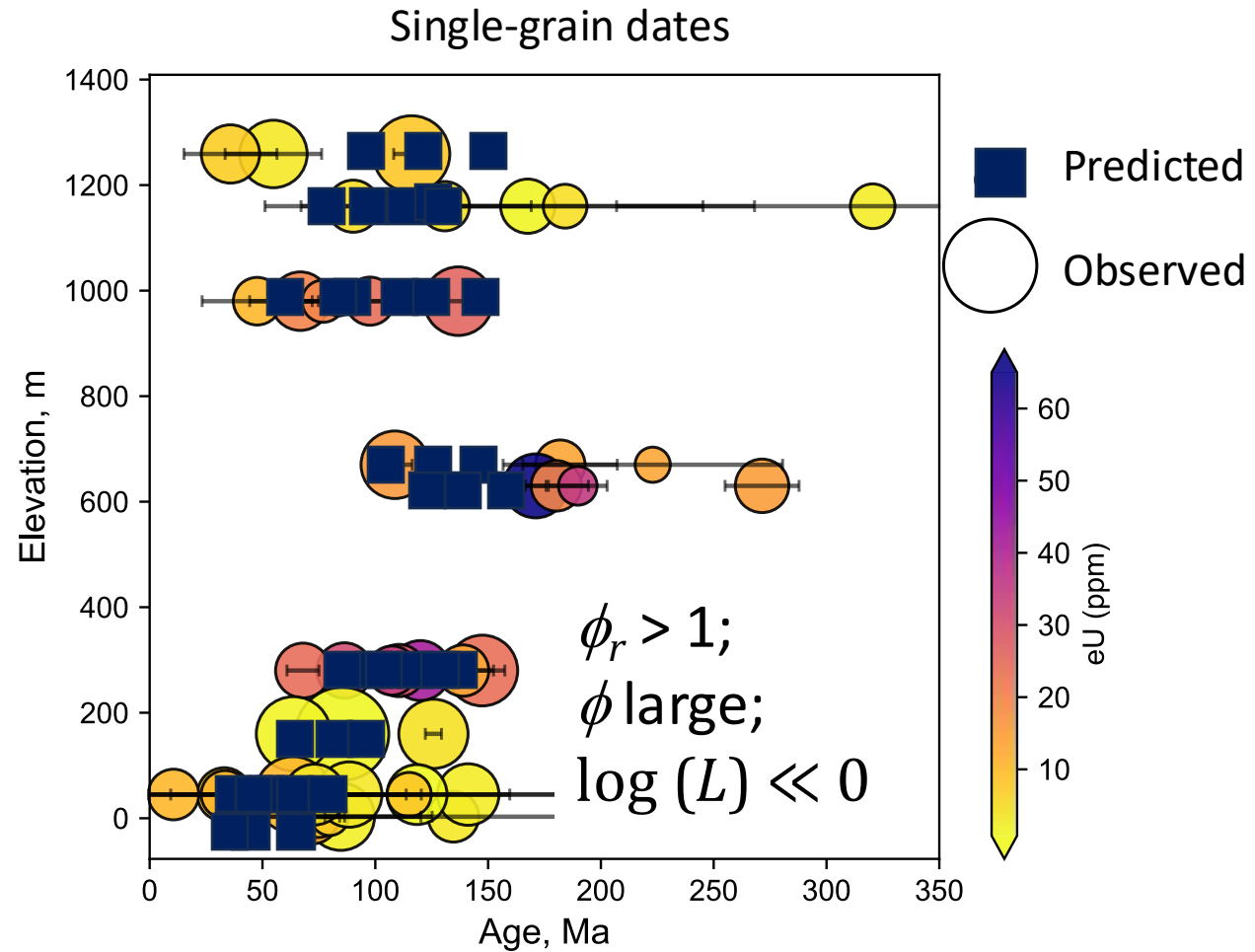
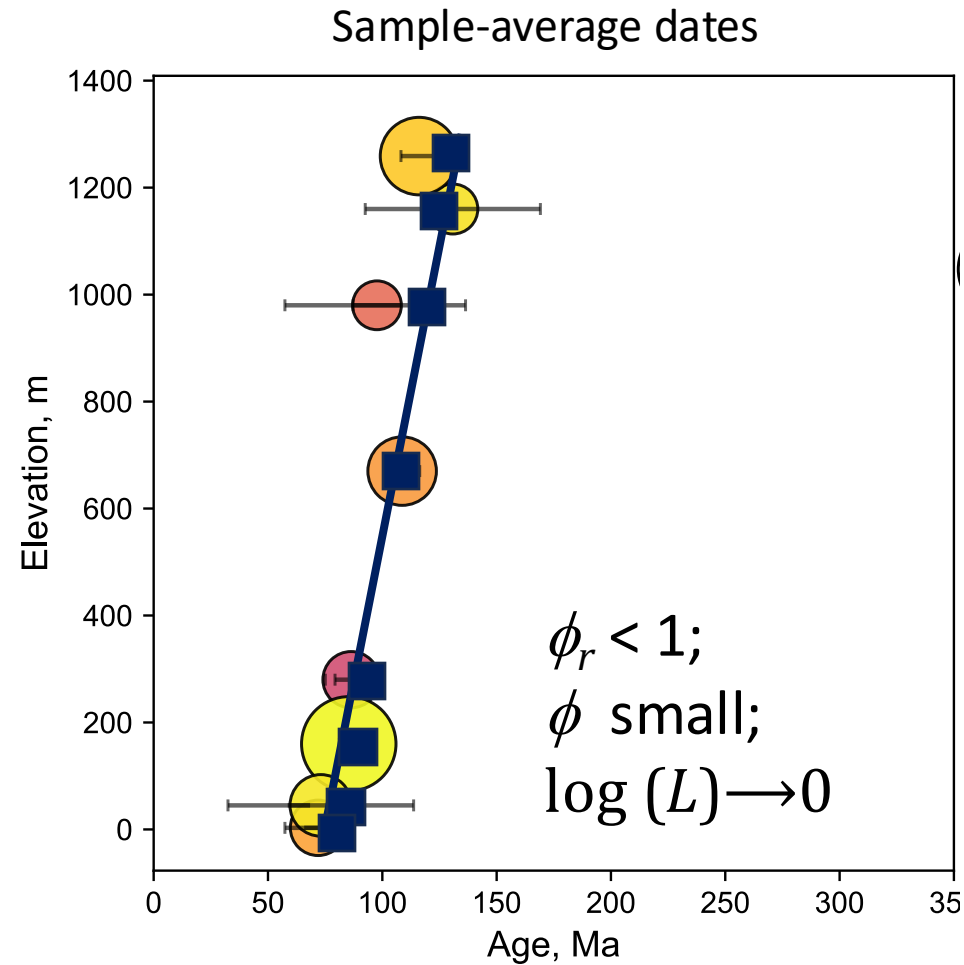
Not such a good  
fit to data?





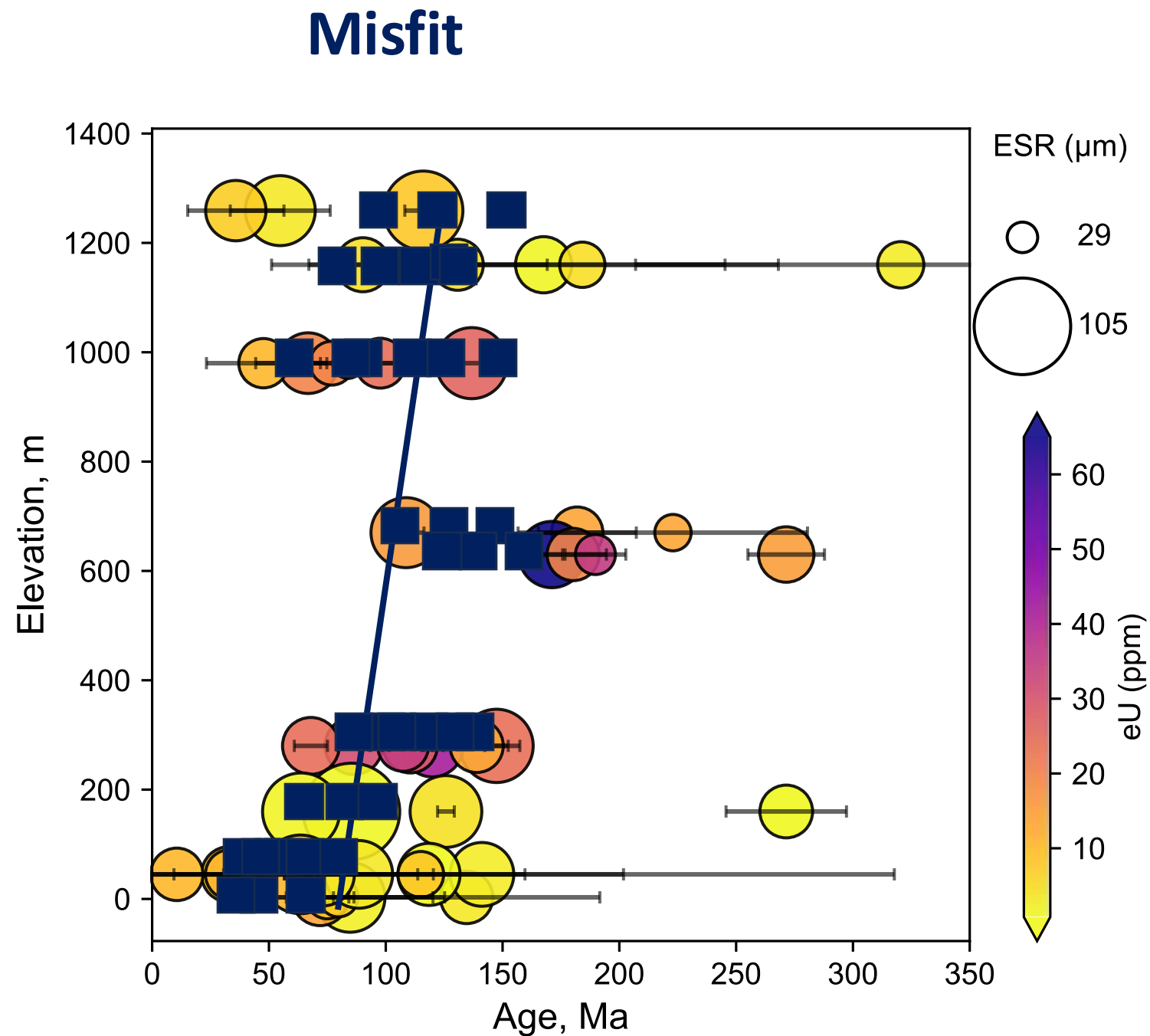
# Misfit

## Working with sample-average versus single-grain data



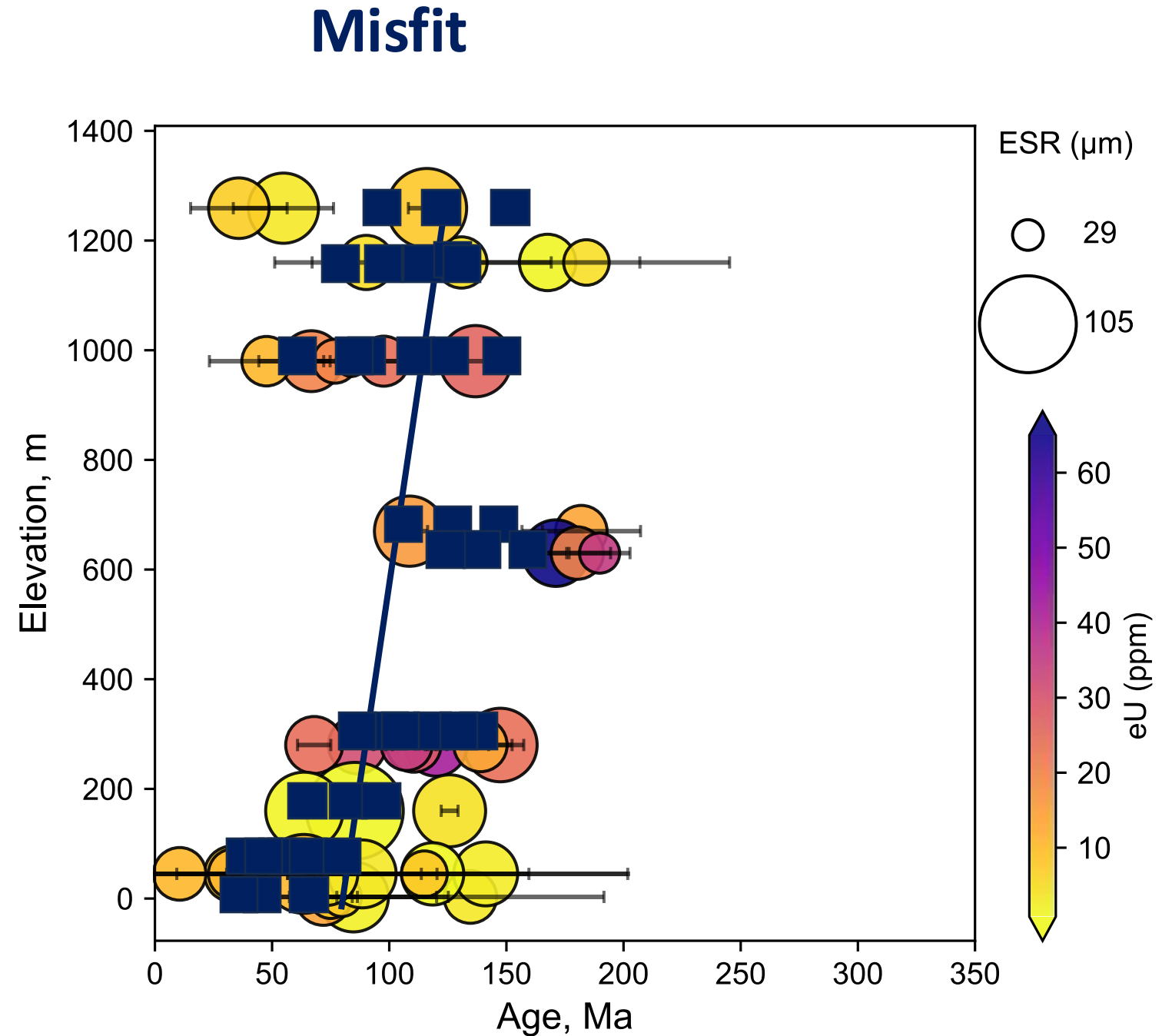
# The role of outliers

$\phi_r > 1$ ;  
 $\phi$  large;  
 $\log(L) \ll 0$



# The role of outliers

$\phi_r > 1$  but smaller;  
 $\phi$  smaller;  
 $\log(L) < 0$



# Inversion in Pecube-GUI

**Forward** model: model  $\rightarrow$  data

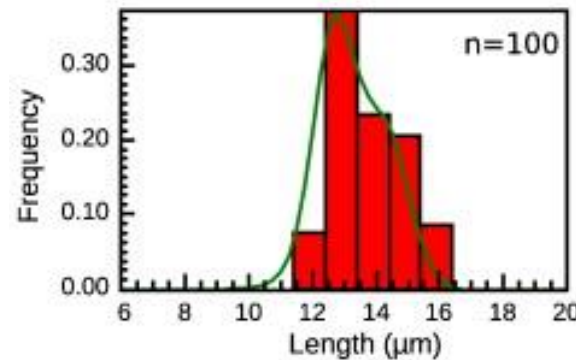
- Scenario testing;
- How well does this model fit my data?

**Inverse** model: data  $\rightarrow$  model

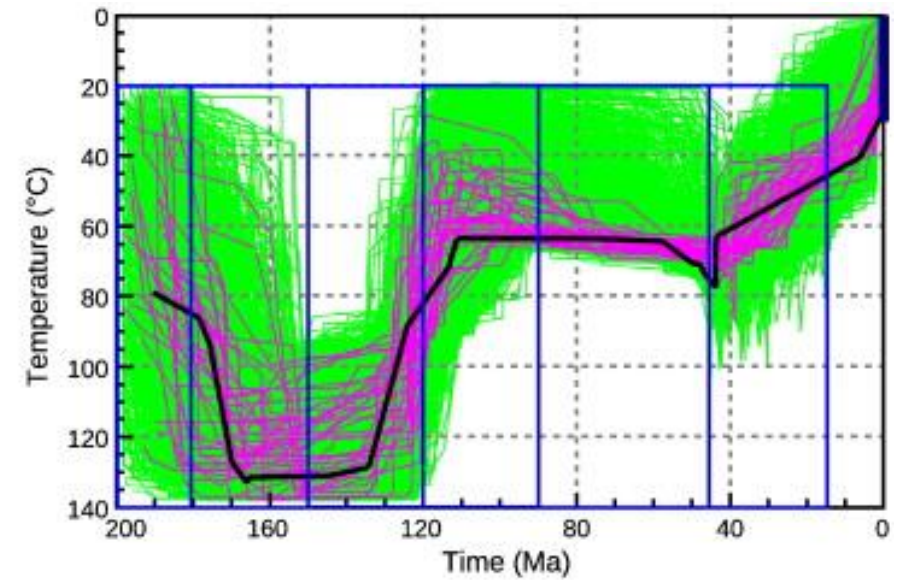
- What does the data say about my model?
- How well can the data resolve contrasting scenarios?

# Thermal history models = inverse models

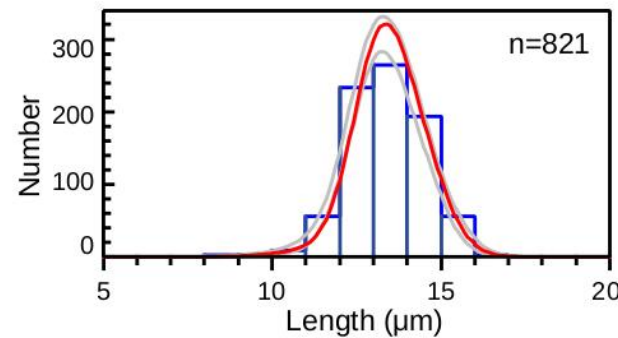
HeFTy



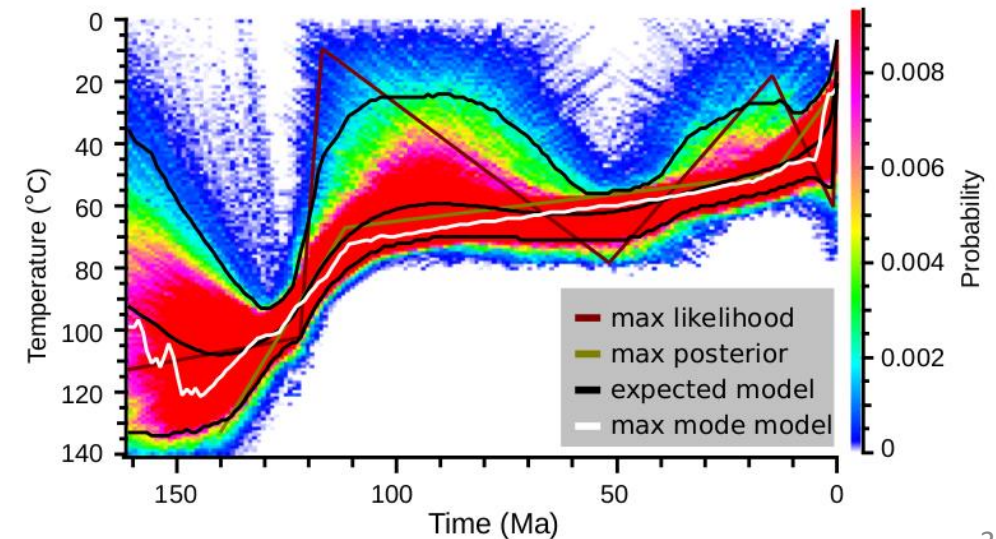
|          | MTL            | AFT age     | AHe age    |
|----------|----------------|-------------|------------|
| Measured | $13.4 \pm 1.0$ | $102 \pm 7$ | $55 \pm 5$ |
| Modelled | $13.3 \pm 1.1$ | 102         | 59         |



QTQt

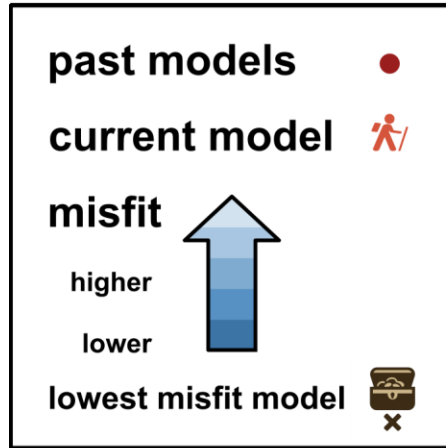


|          | MTL            | AFT age     | AHe age    |
|----------|----------------|-------------|------------|
| Measured | $13.4 \pm 1.2$ | $102 \pm 7$ | $55 \pm 5$ |
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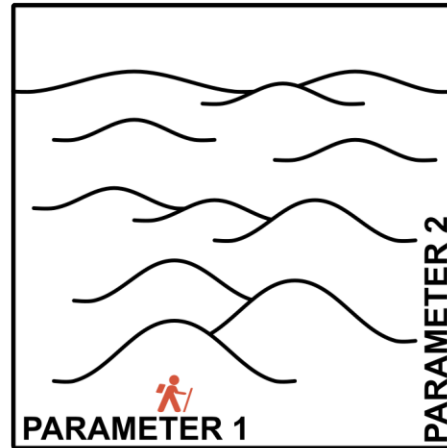


# Inverse modelling strategies

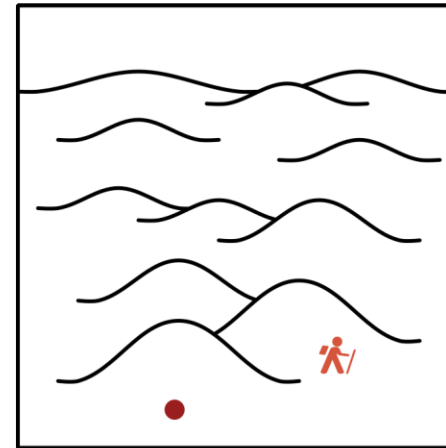
## MONTE CARLO



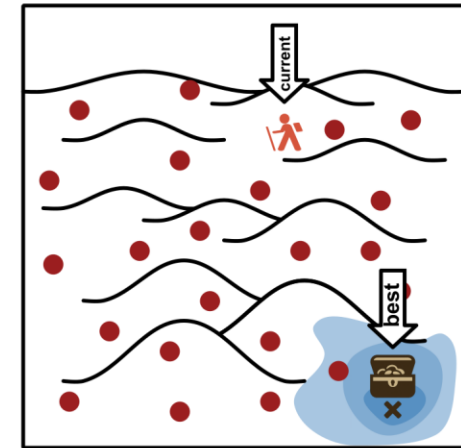
first model randomly chosen  
in parameter space...



...second model...



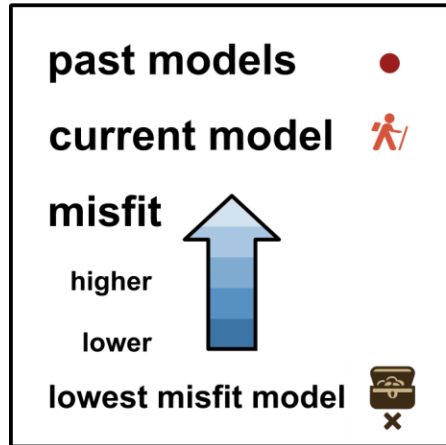
...many models later, random  
new models are investigated  
irrespective of the current  
best (lowest misfit) model.



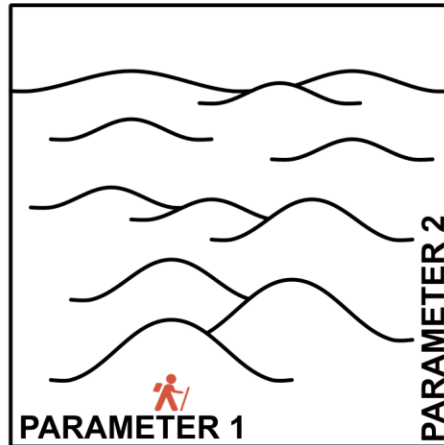


# Inverse modelling strategies

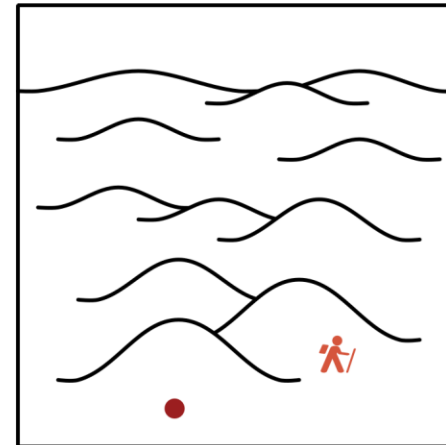
## MONTE CARLO



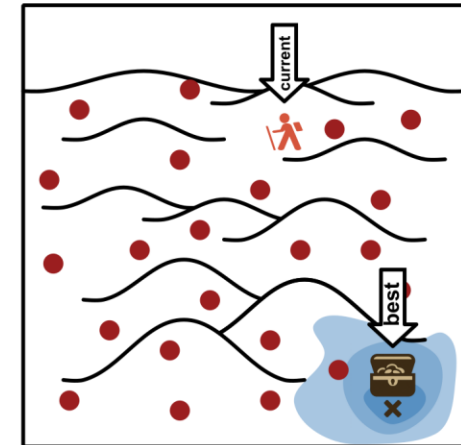
first model randomly chosen  
in parameter space...



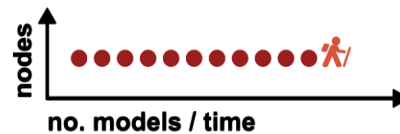
...second model...



...many models later, random  
new models are investigated  
irrespective of the current  
best (lowest misfit) model.



ITERATIONS / COMPUTATION TIME



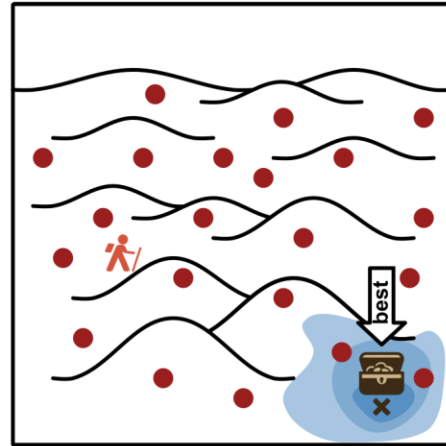
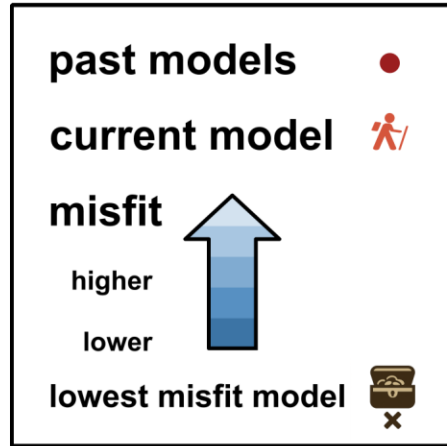
running random models in parallel  
reduces computation time

# Inverse modelling strategies

MONTE CARLO

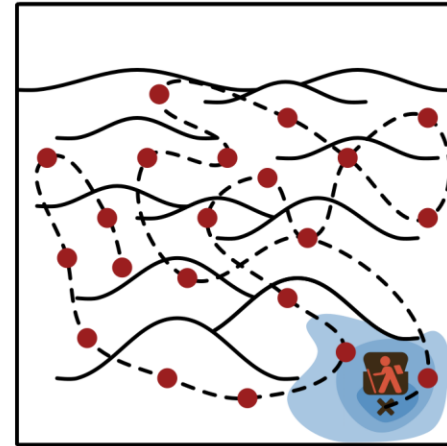
MARKOV CHAIN  
MONTE CARLO

NEIGHBORHOOD  
ALGORITHM



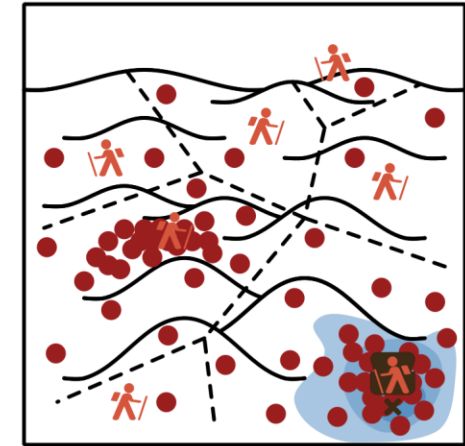
random

e.g., HeFTy



next step based on  
misfit of previous step

e.g., QTQt

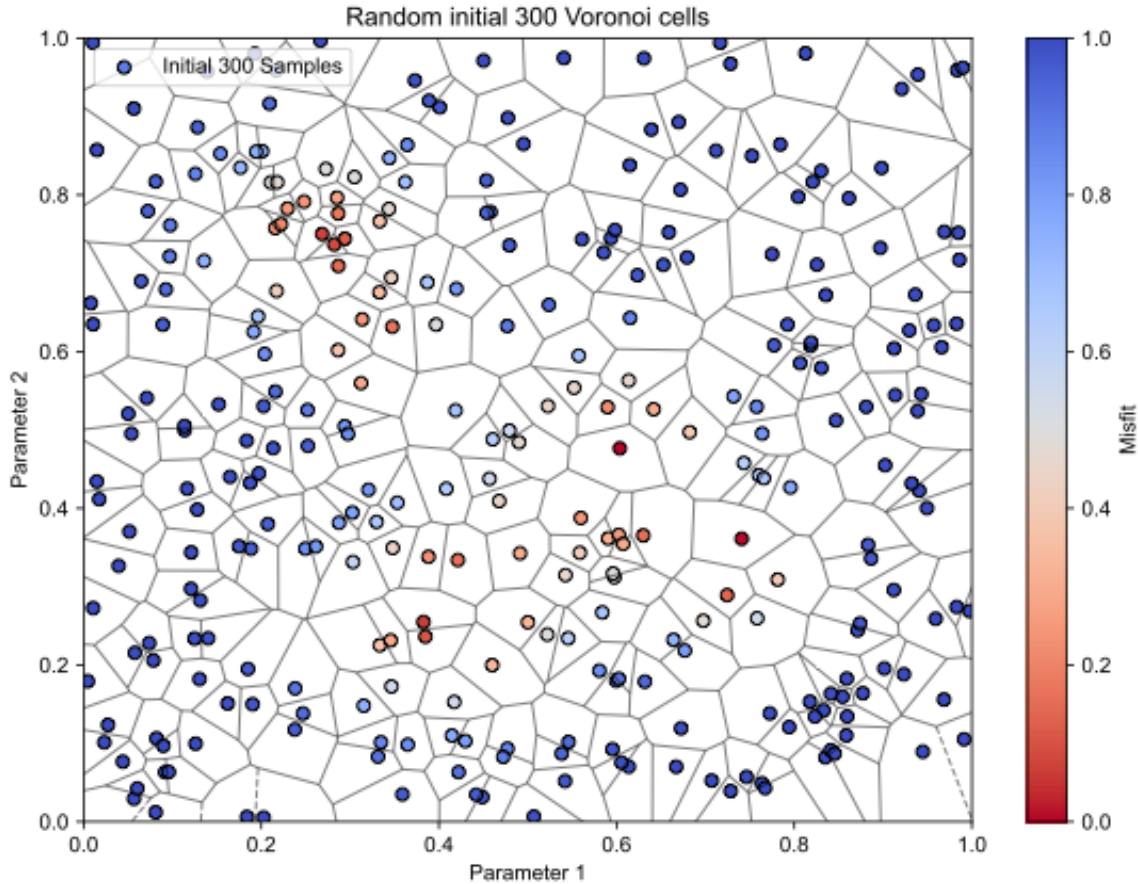


successful neighborhoods  
explored more

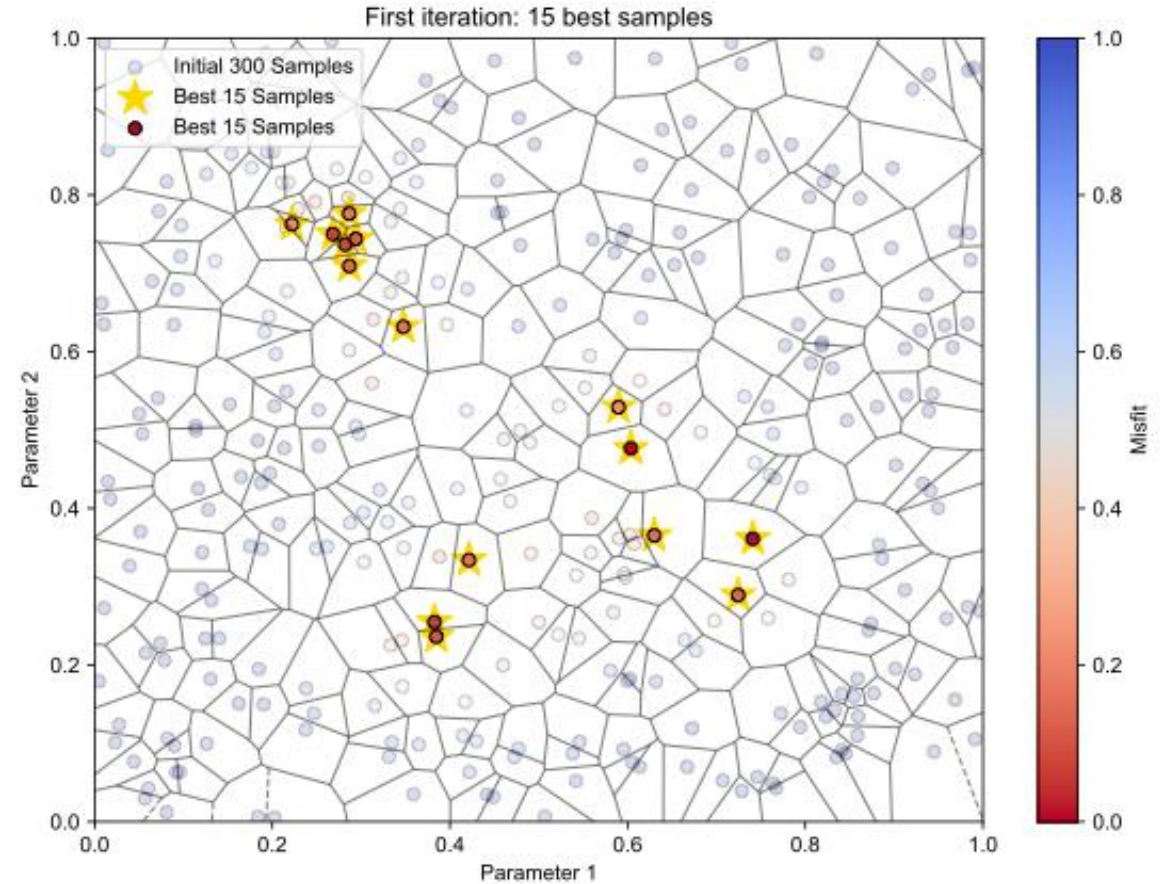
e.g., Pecube

# Inversion in Pecube-GUI: Neighbourhood Algorithm (NA)

Define *nearest neighbours* and divide up model space using *Voronoi cells*

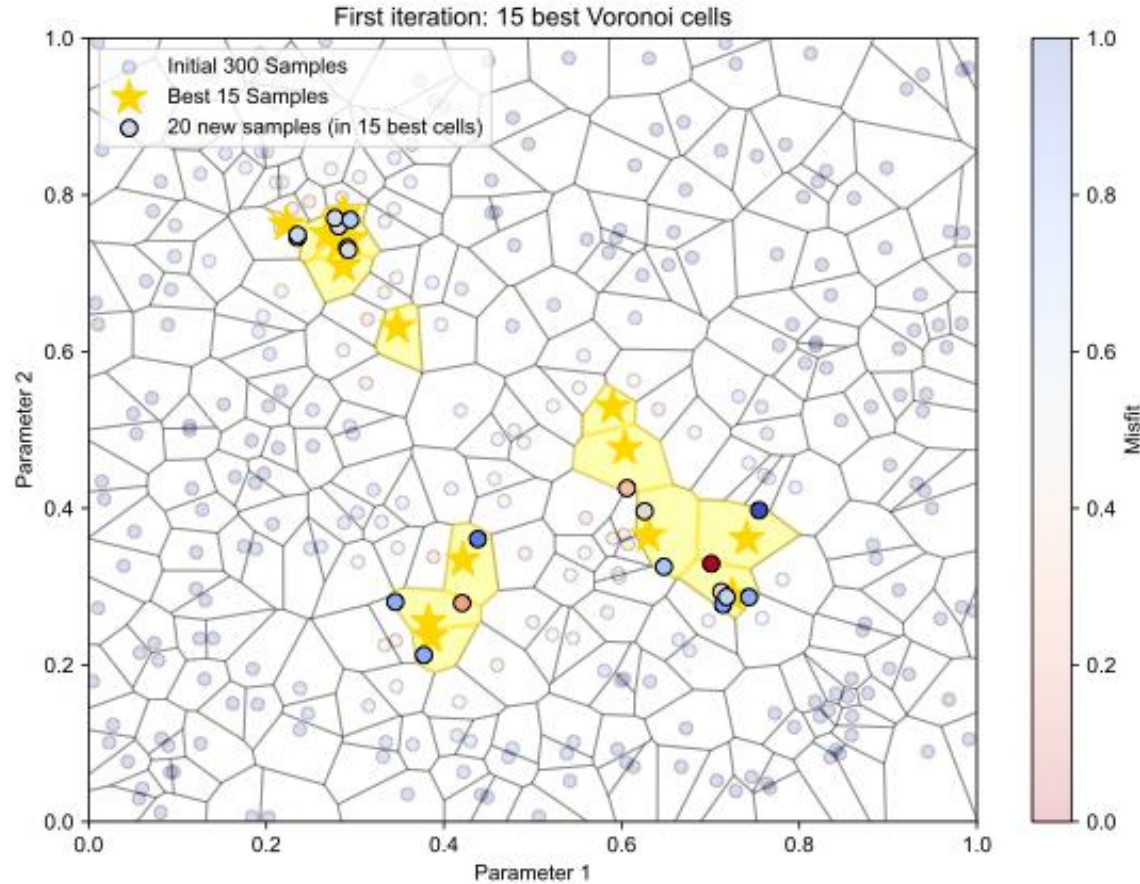


Step 1,  $n$ : search in  $j$  Voronoi cells that contain  $j$  best models from previous step

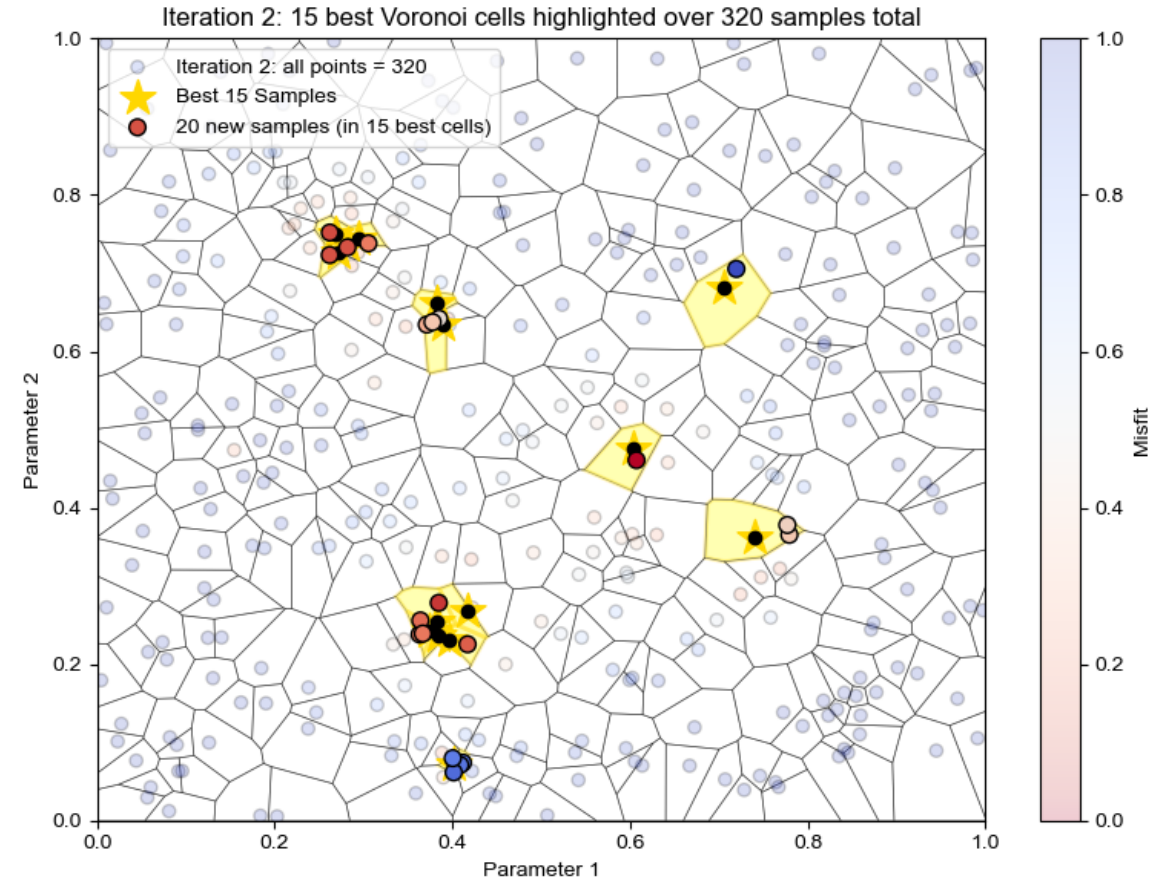


# Inversion in Pecube-GUI: Neighbourhood Algorithm (NA)

Step 1,  $n$ : search in  $j$  Voronoi cells that contain  $j$  best models from previous step



Step 1,  $n$ : search in  $j$  Voronoi cells that contain  $j$  best models from previous step



and so on ...

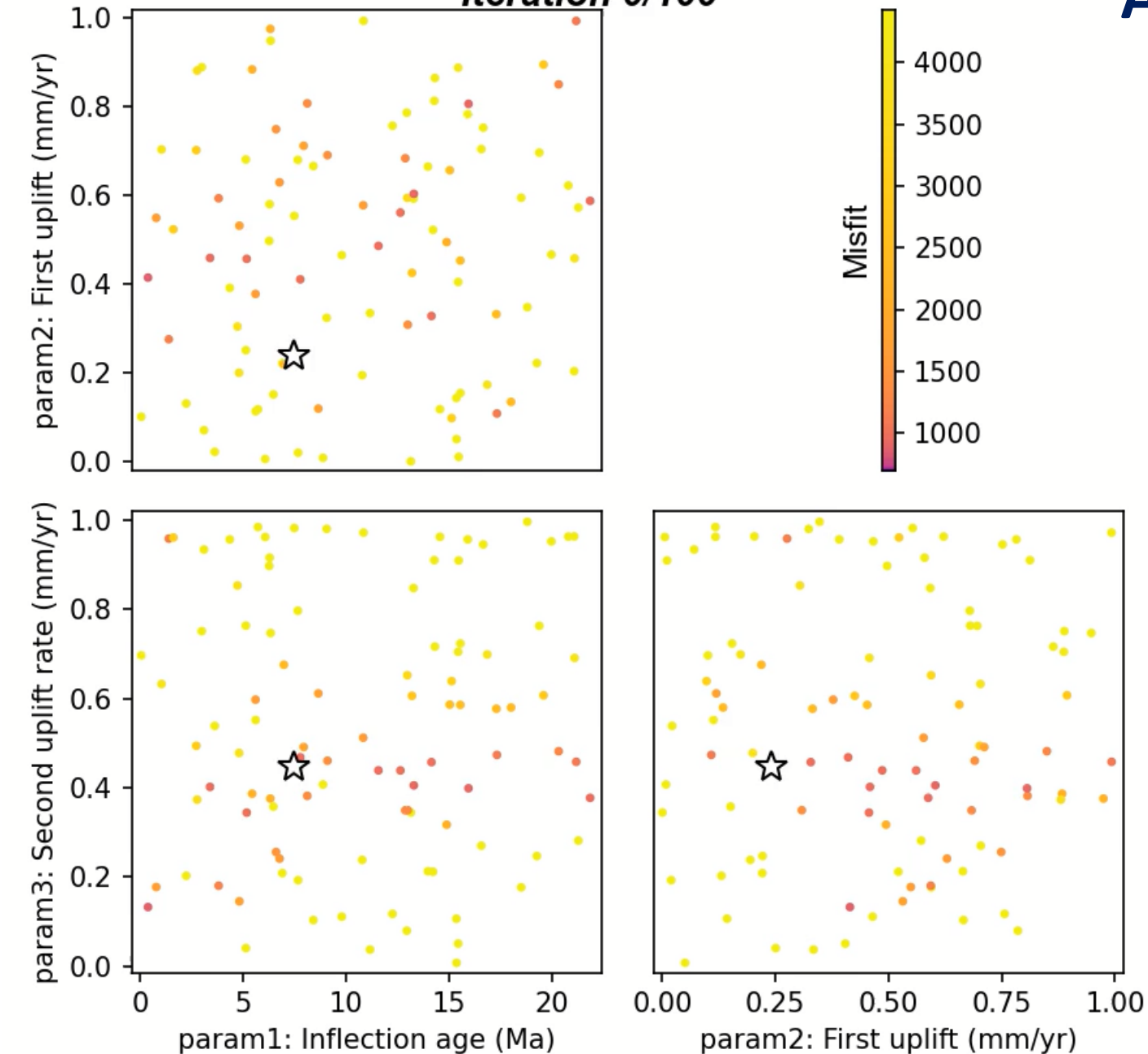


**BAT12: (3 params; 10100 models)**  
**Iteration 0/100**

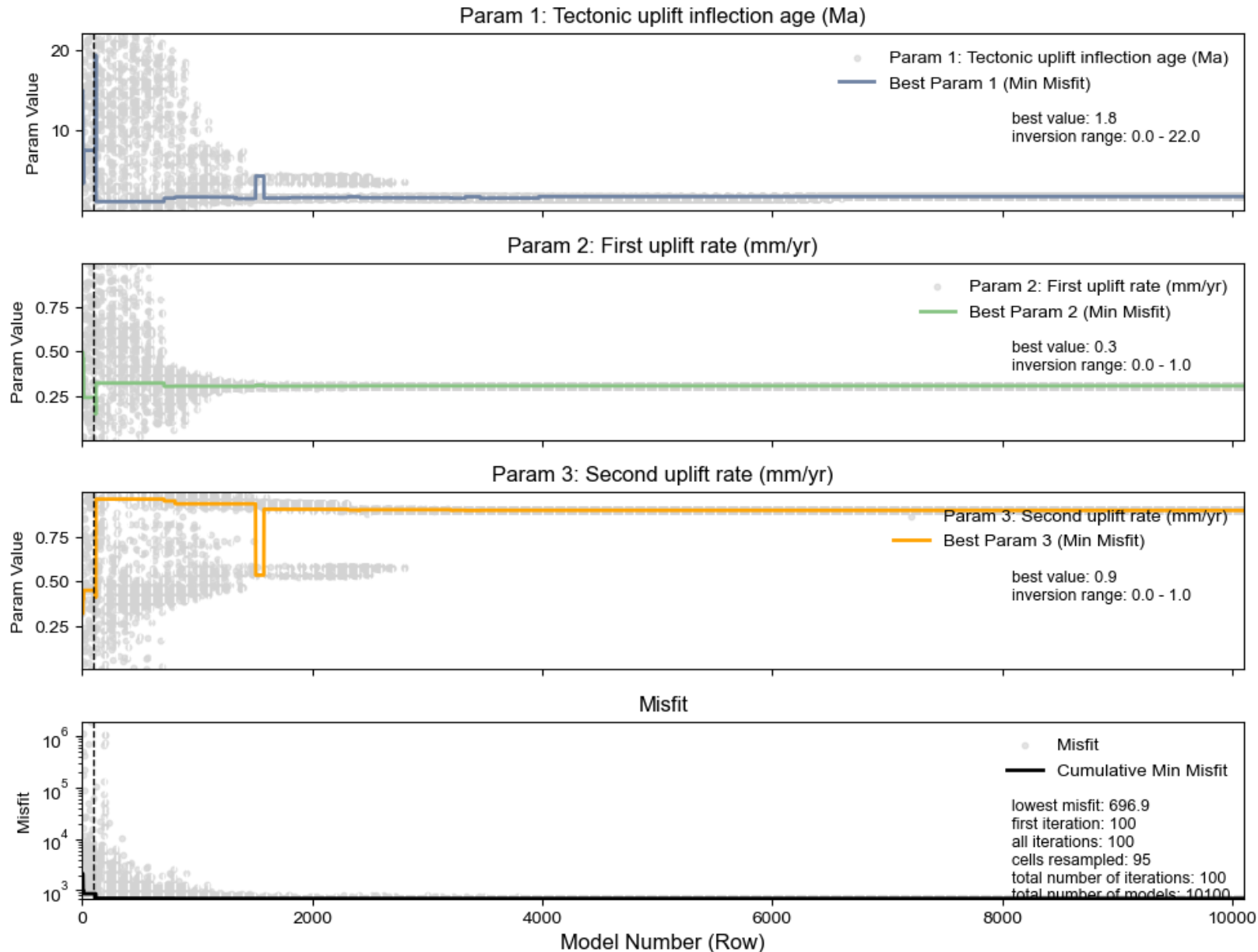
## A simple example

3-parameter model – tectonic scenario only

- Initial rock-uplift rate (param 2)
- Final rock-uplift rate (param 3)
- Time of change (param 1)



# A simple example

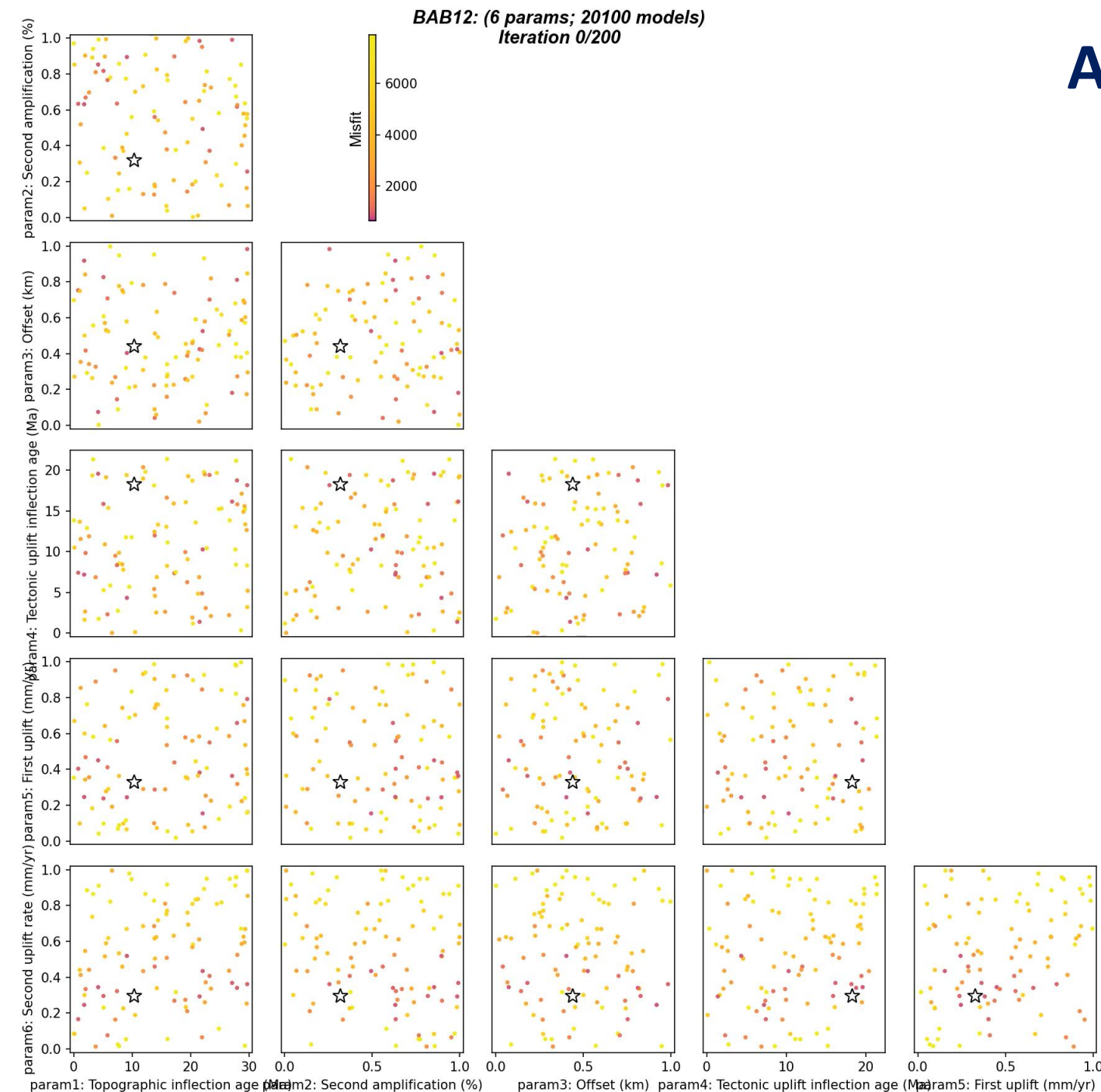




# A more complex example

6-parameter model: tectonic and geomorphic scenario

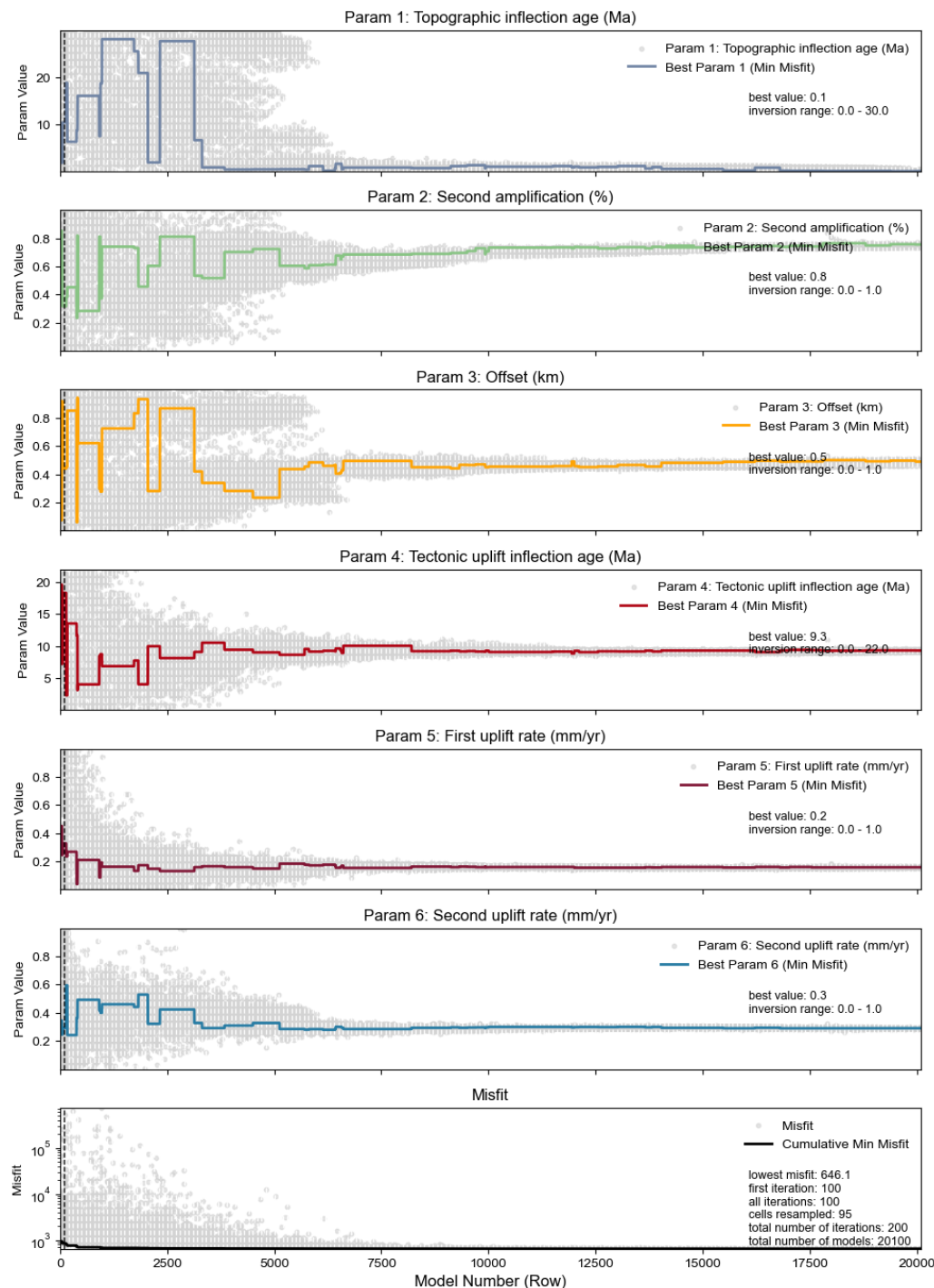
- Initial rock-uplift rate (param 5)
- Final rock-uplift rate (param 6)
- Time of tectonic change (param 4)
- Time of geomorphic change (param 1)
- Initial topographic amplification (param 2)
- Initial topographic offset (param 3)



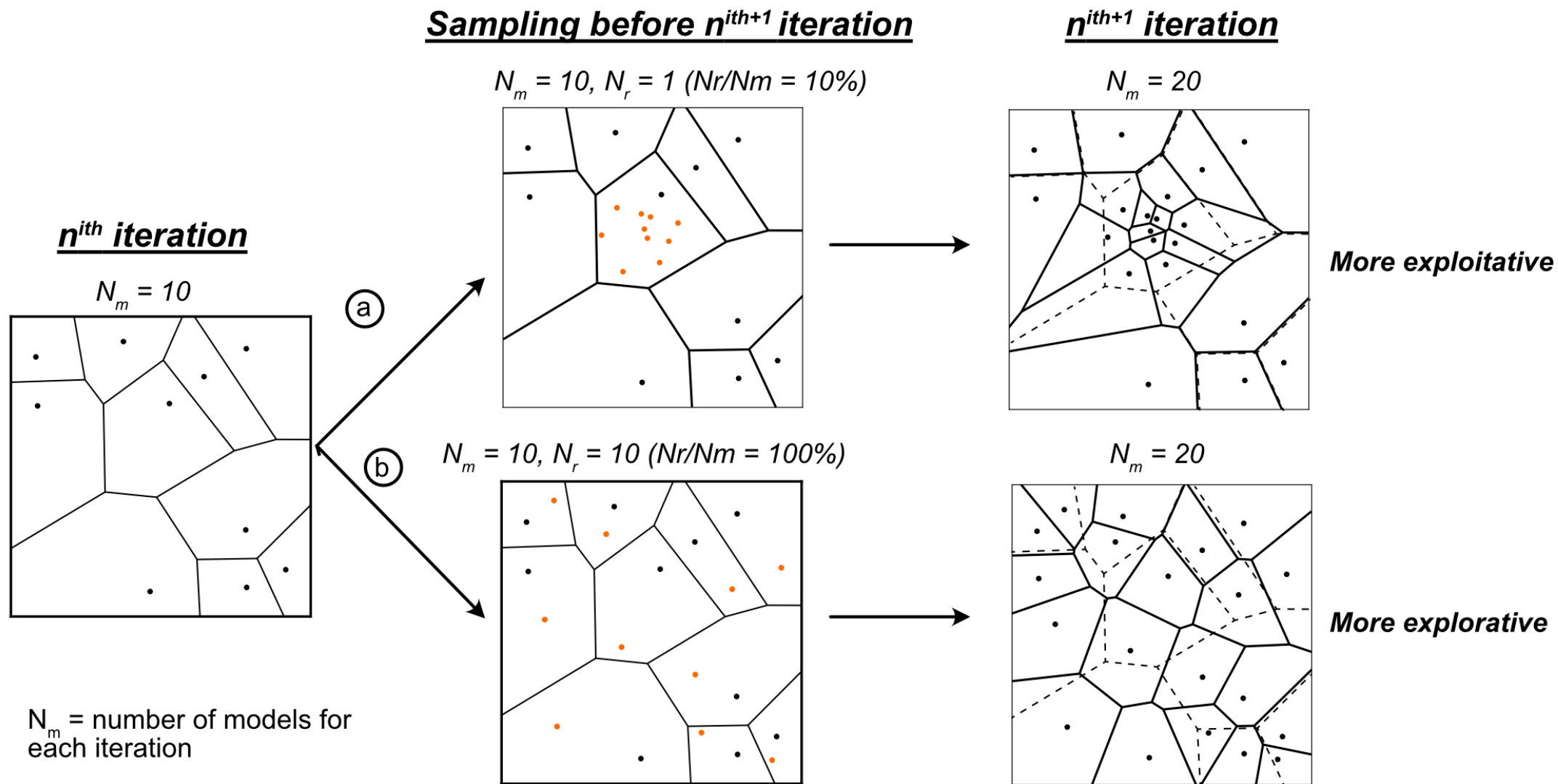
# A more complex example

6-parameter model: tectonic and geomorphic scenario

- Initial rock-uplift rate (param 5)
- Final rock-uplift rate (param 6)
- Time of tectonic change (param 4)
- Time of geomorphic change (param 1)
- Initial topographic amplification (param 2)
- Initial topographic offset (param 3)



# NA inversion strategies



$N_m$  = number of models for each iteration

$N_r$  = number cells to resample (i.e., the  $N_r$  best-models)

After Bernard et al., in prep.

# What does the “best fit” model mean?

After an NA-inversion, you can (it would be better to) run **NA-Bayes** to:

- Get an estimate of resolution;
  - Assess uniqueness
- 
- Initial rock-uplift rate (param 2): well defined
  - Final rock-uplift rate (param 3) ] several minima
  - Time of change (param 1). ] and significant tradeoff

**NA-Bayes** is (currently still) a stand-alone Fortran code but the tools to visualise the predicted parameter PDF's and estimated uncertainties (as shown here) are now available on Zenodo as *Python* and *Matlab* scripts.

See links on: <https://erc-cooler.eu/codes/>

